A Thesis<br>Presented in Partial Fulfillment of the Requirements for the<br>Degree of Master of Science<br>with a<br>Major in Geography<br>in the<br>College of Graduate Studies<br>University of Idaho<br>By<br>JAMES HUNTER MCGEE<br>Major Professor: Grant L. Harley, Ph.D.<br>Committee Members: John T. Abatzoglou, Ph.D., Justin T. Maxwell, Ph.D.<br>Department Administrator: Leslie Baker, Ph.D.

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## AUTHORIZATION TO SUBMIT THESIS

The thesis of James H. McGee, submitted for the degree of Master of Science with a major in Geography and titled, "TESTING THE EFFICACY OF BLUE INTENSITY AS A TEMPERATURE PROXY IN PICEA ENGELMANNII FROM ITS SOUTHERN RANGE LIMIT, NORTHERN NEW MEXICO, USA," has been reviewed in final form. Permission, as indicated by the signatures and dates given below, is now granted to submit final copies to the College of Graduate Studies for approval.

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#### Abstract

Annually resolved temperature proxies are rare for the American Southwest. Recent studies involving the analysis of blue light intensity consistently show an inverse show in inverse relationship between maximum latewood density and blue intensity values. Blue intensity analysis records the amount of blue light absorbed by tracheid cells, thereby quantifying the amount of lignin present in the latewood of the annual rings. This study aims to fill the gap of historical temperature data for the southern range limit of the Sangre de Cristo Mountains of northern New Mexico using annual climate data and blue light intensity analysis of tree rings as a proxy for maximum latewood density analysis. This is done using 27 high-elevation Engelmann spruce (Picea engelmannii Parry ex Engelmann) samples collected from 16 trees located at Wheeler Peak, New Mexico. We also include a test of the efficacy of generating BI data with the CooRecorder density software package. Samples are also measured for minimum earlywood and latewood density, and change in blue intensity values (delta). Results of this study suggest that a statistically significant relationship exists between blue intensity values and maximum annual summer temperature. A warming trend is evident at the turn of the $21^{\text {st }}$ century when observing the delta blue intensity time series data, which is also present in PRISM instrumental temperature data used. Additional research utilizing the methods described in this study can be conducted at similar sites located at high latitudes and alpine environments. Furthermore, a study comparing trends in blue intensity parameters across multiple sites would be valuable. This study contributes to the increased understanding of how blue light intensity can be used as a new and innovative tool within the field of dendroclimatology.


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## CHAPTER ONE

## INTRODUCTION

Anthropogenically induced climate change has resulted in a consistent warming of Earth's climate. Over the past century, Earth's surface temperatures have experienced an increase of roughly $0.6^{\circ} \mathrm{C}$, with the highest rates of increase during 1910 to 1945 and from 1976 to present day (Walther et al., 2002). This trend of increasing global temperatures has scientists searching for innovative methods to determine whether natural climate variability, or pollution onset by human activities is the cause. Because consistently recorded meteorological data is limited to the past century, in order to reconstruct historic temperatures, scientists must identify and utilize temperature-sensitive climate proxies. The use of tree rings as a proxy to reconstruct historic temperature allows scientists to develop more accurate estimations of past climate change (Eschbach et al., 1995). Few temperature-sensitive proxies currently exist across the globe, though they are much more common at high latitudes as well as alpine environments (Wilson and Luckman, 2003; Wilson et al., 2012).

Dendrochronology uses statistical analysis of tree rings to provide useful datasets to various disciplines such as climate science. Dendrochronology aims to assign calendar years to each individual annual tree ring. Because tree rings are unique to the growing conditions during the period they were created, tree rings can produce valuable dendroclimatic data that extends far past instrumental climatic records. Trees can be crossdated using the anatomy of their growth rings because individual specimens that are located in a similar area have similar
growth responses to changing environmental factors, such as precipitation, temperature, and disturbances. The analysis of the annual growth rings of trees provides information for a variety of regional conditions.

Traditionally, dendrochronology consisted of analyzing the observable physical growth patterns of annual rings (Stokes and Smiley, 1968, 1996; Speer, 2010), limiting the proxies in which valuable data is provided. However, recent technological advances in classic dendrochronological methodologies have allowed for individual samples to be scanned at high resolution and analyzed for a number of new climate proxies. The use of specialized software has allowed for the expansion of dendrochronological studies to incorporate other tree ring parameters than tree ring width (TRW) and growth anomalies.

The use of tree ring proxies to reconstruct historical climate variability has been widely documented (Schweingruber and Briffa, 1996; D’Arrigo et al., 2006; Buckley et al., 2018). Though a variety of different proxies can be used to reconstruct climate variability, temperature reconstructions typically rely on expensive specialized equipment to determine the cell densities of individual tree ring samples. Until the turn of the $21^{\text {st }}$ century, dendroclimatological temperature reconstructions were conducted using X-ray densitometry to measure maximum latewood density (MXD) (Eschbach et al., 1995). One of the disadvantages to X-ray densitometry is high operating costs, as it requires specific machinery, software, and trained personnel to process samples. The extensive monetary resources required for this type of analysis is exclusive, as it restricts use to affluent labs; tree ring laboratories that are underfunded may be unable to perform temperature reconstructions using X-ray densitometry (Rydval et al., 2014).

Recent temperature reconstructions have been developed for much of Europe and parts
of the United States, though there are gaps in historical temperature data for much of the American Southwest (Wilson et al., 2016). Existing dendroclimatological studies spanning over the past few decades have suggested that high latitude alpine conifer chronologies of MXD have a much stronger relationship with summer temperatures than TRW-only chronologies (Wilson and Luckman, 2003; Rydval et al., 2014). As with any proxy or tree ring parameter, TRW is useful when reconstructing historic precipitation, though TRW varies based on a number of environmental variables, making it a much less reliable proxy for temperature than cell density. Recently, scientists have been performing an innovative and inexpensive method of calculating the density of tree rings. This new method of densitometry, known as blue intensity (BI) analysis, uses high-resolution scanned images of cores to quantify the blue wavelengths of light that are reflected off of lignified cell tissue.

BI analysis is conducted by measuring several values that quantify the blue wavelengths of light that are reflected and absorbed by lignified cells that make up tree growth rings. The BI parameters utilized in this study are inverted maximum latewood BI and the difference between earlywood BI and latewood BI, also known as delta-blue. Recent studies have suggested that the calculation of delta-blue in statistical testing results in a high positive correlation with the TRW series itself (Bjorklund et al., 2014; Buckley et al., 2018).

Unlike X-ray densitometry, BI analysis is much more affordable, therefore making it much more widely accessible. To process samples using BI analysis, woodworking equipment, a high-resolution scanner, and inexpensive tree ring density software are the only supplies that are needed. By producing similar results to studies utilizing X-ray densitometry, BI temperature reconstructions can be developed globally in small laboratories in far less time, with less effort, at a more reasonable cost. This technological advancement leading to
greater inclusion in tree ring research may increase more widespread interest future in dendroclimatological studies.

This study aims to determine the effectiveness of performing BI analysis at the southern range limit of Engelmann spruce (Picea engelmannii Parry ex Engelmann) and to contribute to the historical temperature data available for the American Southwest. The samples used for this study were collected from the southern extent of the Sangre de Cristo Mountains, a remote mountain range that extends from central Colorado to northern New Mexico. This mountain range comprises the southernmost sub range of the Rocky Mountains, which extends through much of western Canada and the United States. The Sangre de Cristo Mountains, in addition to being the southern range of the Rocky Mountains, are also the location of the southernmost growth limit of Engelmann spruce.

### 1.1 Research Questions

This study aims to answer the following questions:

- Can total ring width of Engelmann spruce be used as a predictor of historical temperature for its southern range limit in the Sangre de Cristo Mountains, NM?
- Does blue intensity analysis reveal significant correlations with instrumental temperature in Engelmann spruce growth rings at its southern range limit?


## CHAPTER TWO

## LITERATURE REVIEW

### 2.1 Temperature Reconstructions Using Tree Rings

Anthropogenic climate change has quickly become an influential topic of discussion for Earth scientists. Recent studies observing instrumental climate records have suggested that Northwestern North America has one of the highest increasing rates of recent temperatures globally (Hartmann and Welder, 2005). Instrumental climate data has a finite timescale, thus, there has been a growing interest in the use of proxy climate records to reconstruct historical fluctuations in climate. Early identification of the potential drivers of global climate change can allow for the development of policies to reduce human influence on climatic alterations.

Proxy records, such as tree rings, are interdisciplinary in nature, spanning across biological and geological sciences, and varying in spatial and temporal resolutions (Stokes and Smiley, 1968). TRW is used as a high-resolution palaeoclimate proxy for historical climate variation (Wilson et al., 2016). Due to its relatively short timescale in relation to other climate proxies such as pollen, TRW is not a useful low-resolution climate proxy. A vast majority of the temperature reconstructions for the Northern Hemisphere are developed from TRW data (D’Arrigo, et al., 2006; Frank, et al., 2007; Wilson, et al., 2016), though maximum latewood density has been shown to be a better recorder of annual summer temperatures (Wilson, et al., 2016).

Since the 1980s, notable dendroclimatologists such as Fritz Schweingruber, Keith Briffa, and Jan Esper have been developing a global network of tree ring density data (Briffa and Schweingruber, 1988, 1992; Schweingruber and Briffa, 1996; Esper et al., 2002). The
studies these scientists have conducted observe fluctuations in the MXD of growth rings of tree species that express variations in growth trends related to temperature. Specifically, studies observing MXD as a temperature proxy have found that a positive relationship exists between MXD and maximum annual summer temperature (Briffa et al., 2001, 2002). Additionally, they use specialized equipment and methodology that is both time-consuming and expensive, limiting it to laboratories with ample resources. A rise in the demand for a less costly method to measure tree ring density has surfaced in an attempt to construct a more complete record of historic temperature variability globally.

### 2.2 Development of Blue Intensity Analysis

Due to the high cost associated with performing X-ray densitometry to reconstruct past temperatures, scientists have been interested in developing a more practical method of determining the cell density of tree rings. The absorption of ultraviolet light by lignified cell tissue, known as ultraviolet microscopy, has long been utilized as a tree ring climate proxy (Lange 1954; Fukazawa 1992). More recently, it has been discovered that with the use of a basic optical scanner and specialized computer software, the reflectivity of certain wavelengths of the visible light spectrum can be quantified (McCarrol et al., 2002; Björklund et al., 2014). Lignified cells readily absorb wavelengths of light in the blue spectrum, thus quantifying the amount of blue light that is reflected off of tracheid cells can provide insight about the amount of lignin that is present in dense latewood bands of cell tissue. The cells of latewood band of tree rings are typically darker than its earlywood counterparts, meaning the latewood will absorb more blue light than it reflects.

Proxies that use cell density to observe past temperatures have been successfully utilized over the past several decades (Briffa et al., 1992, 2001; Schneider et al., 2015; Wilson
et al., 2016a). Maximum latewood density of annual growth rings has been recognized as a recorder of past temperature variability (Büntgen et al., 2006; Rydval et al., 2014). Several studies have utilized X-ray densitometry to measure MXD, which has been shown to have a positive relationship with maximum annual summer temperatures (McCarroll et al., 2002; Wilson and Luckman, 2003; Wilson et al., 2016a; Rydval et al., 2017). Many of these existing studies have been conducted across European and Asian countries, as well as Canada, though few have been performed in the United States. This has resulted in a lack of tree ring density data for much of North and South America, likely from the high operation costs associated with X-ray densitometry (Rydval et al., 2014).

With technological advancements in the field of dendroclimatology, the absorption of visible spectrums of light by lignified tracheid cell tissue can be quantified and compared to global temperature data. Wilson (2012) found that this value of absorption and reflectivity, known as blue intensity, has an inverse relationship with past maximum summer temperatures. By producing high-resolution scans of tree ring samples and generating the various BI parameters offered by dendro-specific software such as CooRecorder (Rydval et al., 2014) or WinDendro (Campbell et al., 2011), scientists can have a good understanding of the density associated with various sections of cells making up annual tree growth rings. Additionally, dendro-specific software allows the user to generate both BI data as well as TRW data simultaneously, making the production cost of BI time series data similar to the cost associated with the development of tree ring chronologies.

Though the BI method of generating density data is far more practical in cost and methodology that X-ray densitometry, there are some limiting factors. BI is still a new practice, with the first studies surfacing in the early 2000s (McCarroll et al., 2002; Wilson and

Luckman, 2003). The relatively young age of this type of research has resulted in a lack of BI information in several regions across the globe, most notably North America (Wilson et al., 2016a). Additionally, as with X-ray densitometry (Briffa et al., 1992), useful BI data can only be produced from temperature-sensitive conifers at mid to high latitudes and alpine environments (Wilson et al., 2017).

### 2.3 Engelmann Spruce

Engelmann spruce, named for physician and botanist George Engelmann, are members of the Pinaceae family, endemic to North America. Their habit is typically characterized as a larger tree, growing upwards of 60 meters tall, with a relatively dense crown dense that forms a narrow cone or spire-like shape. Engelmann spruce branches tend to exhibit horizontal spreading. The lower branches are generally persistent, as it is not a strongly self-pruning species. The bark is red to purplish-brown in color and thin and scaly. Needles are evergreen, borne singly from all sides of stout, yellow-brown twigs. Needle length ranges from $1.6-3.0 \mathrm{~cm}$ long. Needles are 4-angled, stiff, sharply pointed on the ends, and blue-green in color. Seed cones are violet to deep purple in color. Upon ripening, cones turn dull-brown, ellipsoid, and pendent, with lengths ranging from $3.0-6.0 \mathrm{~cm}$. The cone scales are relatively small, papery, and flexible, and generally remain intact after cones drop off the tree (Alexander and Shepperd 1990).

Engelmann spruce are endemically distributed from Alberta and British Columbia to the north, and southward through Nevada, Utah, and Colorado, into Arizona and New Mexico. The Sangre de Cristo Mountains are at the southernmost extent of its range. Populations further south such as populations in the Chiricahua Mountains and those down in northern Mexico are considered a subspecies (ssp. mexicana). Engelmann spruce occur in
montane and subalpine forests. Engelmann spruce and subalpine fir (Abies lasiocarpa) form one of the most common forest associations in the Rocky Mountains. With an elevation range anywhere from 1,000-3,000 meters, these trees can occur as stunted, twisted individuals at timberline, or even co-occur down into the fir-aspen belt on moist, north facing slopes and in canyons (Alexander and Shepperd 1990).

The strong shade tolerance of Engelmann spruce allows it to occur both as a persistent long-lived seral species and as a major climax species (Aplet et al. 1988; Alexander and Shepperd 1990). Engelmann spruce will grow steadily for 300 years, long after the growth of most associated tree species slows down (Aplet et al. 1988). Dominant spruces often range from 250-450 years old, and individuals 500-700 years old have been documented in The Old List (http://www.rmtrr.org/oldlist.htm). The oldest recorded specimen for this species was sampled by Brown et al. (1995) with an age of 911 years.

While open-growth trees may begin producing seed crops as early as 15 years, the best seed production for Engelmann spruce occurs between 150 and 250 years. Significant seed crops are generally born every $2-5$ years. Germination and establishment in typically occur on duff, litter, humus, decaying wood, and mounds of mineral soil upturned by wind thrown trees (Knapp and Smith 1982). Engelmann spruce seedlings do not readily establish in totally open conditions (Knapp and Smith 1982). At high elevations, 40 to 60 percent of full shade is most favorable for seedling establishment (Alexander and Shepperd 1990). Because of their slow initial root penetration and extreme sensitivity to heat in the succulent stage, spruce seedlings are often largely killed due to drought and heat girdling in their first year. Drought losses
continue to be significant for the first five years of growth for Engelmann spruce seedlings (Alexander and Shepperd 1990).

Because of having a shallow root system, Engelmann spruce is highly susceptible to windthrow. Downed wood from windthrow also makes a site vulnerable to attack from the spruce beetle, which has caused severe damage in recent years. The western spruce budworm is another potentially damaging insect that attacks both Engelmann spruce and subalpine fir. Complete removal of a spruce-fir stand by fire or logging results in such drastic environmental changes that spruce and fir are usually replaced by lodgepole pine, aspen, or shrub and grass communities (Alexander and Shepperd 1990).

## CHAPTER THREE

# TESTING THE EFFICACY OF BLUE INTENSITY AS A TEMPERATURE PROXY IN PICEA ENGELMANNII FROM ITS SOUTHERN RANGE LIMIT, NORTHERN NEW MEXICO, USA 

### 3.1 Introduction

Increasing global temperatures are continually exacerbated by anthropogenic climate change. Environmental proxies that record fluctuations in historical temperature are crucial to reconstructing and understanding past climates. Currently, climate proxies that are sensitive to temperature, and therefore useful for historic temperature reconstructions, are rare in comparison to proxies used to reconstruct precipitation, drought, and other climate variables. Temperature-sensitive proxies that are currently utilized are more prevalent in mid to high latitudes and in high elevation environments (Campbell et al., 2011; Wilson et al., 2016a).

The study of tree rings to document and place some chronological context to, environmental conditions, ecological processes, and disturbances is known as dendrochronology. Dendrochronology has a number of subfields that have different foci, from dendropyrochronology, which observes historic fire occurrence, to dendroclimatology, which reconstructs historic weather and climate variables. Calendar years are assigned to individual tree growth rings, and the growth structure of each ring is analyzed to determine the environmental conditions present during the time of growth. Dendrochronological analysis can provide a detailed historical record climate that is much longer than the instrumental climate records that are available.

Dendrochronological studies provide detailed reconstructions of a number of regional historical climate variables. Studies using tree rings as a proxy for climate traditionally use visible growth trends such as total ring width (TRW) to reconstruct climate. Programs used to analyze tree ring data can observe a number of additional of additional variables. The datasets from these proxies are extracted from high-resolution scanned images of samples. Dendrospecific software such as Windendro and CooRecorder use these scanned images to delineate tree rings, crossdate samples, measure total ring width, analyze earlywood and latewood, and perform density analysis.

One of the most widely used annually-resolved proxies for climate reconstructions is dendrochronology (Esper et al., 2004). Proxies that reconstruct temperature are rare, though trees of certain species in specific environments can be used to perform these reconstructions. X-ray densitometry is a method of calculating maximum latewood density (MXD), a biological indicator of temperature (Polge 1966; Schweingruber 1988). The density of latewood bands of cells is shown to fluctuate with varying temperatures, though traditional methods of densitometry are rather expensive and required specialized equipment and training. Blue intensity (BI) analysis is a much more widely accessible and affordable method of reconstructing regional temperatures. BI records the amount of blue light that is absorbed by tracheid cells, which provides a numerical reflectivity value that expresses the lignin presence of the latewood of tree growth rings.

To date, there are still many data gaps for temperature reconstructions globally. Temperature reconstructions currently exist for many locations across Europe, as well as a few places in North America (Wilson et al., 2016; Figure 3.1). The American Southwest is one region where no temperature reconstructions have been conducted. Additionally, on a
global scale, no temperature reconstructions have been successfully completed as far south as northern New Mexico. This study aims to determine the efficacy of performing BI analysis at high-elevation mixed conifer forest ecosystems in the Sangre de Cristo Mountains of the American Southwest.


Figure 3.1: Map of the N-TREND network of tree-ring reconstructions depicting lack of North American temperature reconstruction data used in Wilson et al (2016).

### 3.2 Study Site

This study took place in the Sangre de Cristo Mountain range, which compose the southernmost region of the Rocky Mountains and extend from southern Colorado to northern New Mexico. The Carson National Forest, which lies within the boundaries of the Sangre de Cristo Mountains in north-central New Mexico, encompasses the field site, Wheeler Peak, located at $36^{\circ} 33^{\prime} 25^{\prime \prime} \mathrm{N} 105^{\circ} 25^{\prime} 01^{\prime \prime} \mathrm{W}$. Wheeler Peak National Wilderness is located just
outside of Taos, New Mexico in the Taos subrange of the Sangre de Cristo Mountains. Wheeler Peak National Wilderness encompasses an area of approximately 8,100 ha and ranges in elevation from 1828 m to 4011 m . The summit of Wheeler Peak sits at just over 4011 m in elevation, making it the highest point of the State of New Mexico.

In Taos County, NM, the monthly annual temperatures range from a low of $-11.8^{\circ} \mathrm{C}$ in January to a high of $30.4^{\circ} \mathrm{C}$ in July, with monthly annual precipitation ranging from 13.97 mm in February to 52.3 mm in August (NCDC 2018). The variable elevation, slope aspects, and soil types of Carson National Forest has resulted in a landscape with a mosaic of varying vegetation types. The Sangre de Cristos are a particularly structurally complex range, with igneous, metamorphic, and sedimentary rocks visibly exposed at high elevation sites (Baker 1973). The geology from the Wheeler Peak Wilderness Area primarily consists of Precambrian granite, gneiss, and migmatite (Clark and Read, 1972). Of the area composing the Carson National Forest, 87 percent fall into the classification of forested lands (Menlove 2004). This forested land is majorly composed of pinyon-juniper (Pinus edulis; Juniperus scopulorum), ponderosa pine (Pinus ponderosa), and other mixed conifer woodlands (Menlove 2004). Though only a small portion ( $\sim 4 \%$ ) of the Carson National Forest is made up of Engelmann spruce trees, this region comprises the southern range limit of the species, making the specimens at this field site sensitive to environmental fluctuations.


Figure 3.1: Map of the study site (red circle) in the Carson National Forest of the Sangre de Cristo Mountains, New Mexico. Generated using ArcGIS.


Figure 3.2: Map showing distribution Engelmann spruce across North America. Taken from USGS on 28 March 2018.


Figure 3.3: Photograph taken from the field site at Wheeler Peak, NM by Trevis Matheus, August, 2016.

### 3.3 Methodology

### 3.3.1 Field Sampling

Increment cores were collected from a variety of ages and sizes of Engelmann spruce, further on referred to by the species code: PIEN, at the high elevation ( 3500 to 4000 m ) mixed-conifer site. Cores were taken at breast height using a 5 mm diameter increment borer. A total of 27 increment cores were randomly sampled from 16 PIEN located at the tree line. A minimum of two cores were taken from each tree, parallel to the contour of the slope to reduce abnormalities in ring growth and to increase the likelihood of sampling as many rings between the bark and pith as possible (Tucker 1979; Speer 2010). Upon extraction, each core was labeled and placed into protective packaging for transport. Additionally, at each specimen GPS coordinates were recorded for each individual tree that was sampled.


Figure 3.4: Image taken as example of increment core extraction technique

### 3.3.2 Sample Preparation and Laboratory Analysis

Increment cores were prepared for analysis using the standard procedures described by Stokes and Smiley (1968; 1996). Samples were air-dried for a minimum of 24 hours. The cores were then positioned to orient the tracheid cells vertically and secured to 5.0 mm wooden mounts using Elmer's multi-purpose glue and fasteners. The mounted cores were then allowed to dry for additional 24 hours. Once the samples were dried and mounted, each sample was progressively sanded using a rotating belt sander (80 grit, 120 grit, 220 grit, 320 grit, 400 grit, 800 grit). If samples contained any noticeable scratches or blemishes from mechanical sanding, individual samples were sanded by hand during analysis using 1200-grit
sandpaper. The final hand polish removed any imperfections and provided a clearer scanned image.

Prepared increment cores were then scanned into dendro imaging software, CooRecorder. CooRecorder required initial scanner calibration to ensure accuracy of generated BI values. The color intensity values were calibrated using an IT8.7/2 calibration card developed by LaserSoft Imaging coupled with EPSON Scan 2 software and an EPSON XL 12000 scanner. The calibration card was first scanned using 1200 dpi (dots per inch) resolution and 48-bit color parameters. Using CooRecorder's calibration function and a colormetric data file calibrated specifically to the R170419 calibration card that was used, the initial image of the calibration card was visually examined to ensure that each frame was aligned with its appropriate color (Figure 3.5).


Figure 3.5: Screenshot from CooRecorder showing blue color intensity calibration card with colorimetric data file.

At this point, CooRecorder defined three corresponding color values for each frame, prompting the user to save the "calibration points" file. Coorecorder then generated a scatterplot from the calibration points file and plots them against a line of perfectly calibrated color values. If the image is calibrated correctly, most of the color intensity points will be positioned on or close to the calibrated line. If color intensity values vary from the calibration line, the image is not calibrated correctly, and depending on the value that is misplaced, lower values (darker colors) or higher values (lighter colors) may express discoloration. By generating a reasonably good calibration curve through these methods, additional calibration is not necessary, though more accurate calibration is available by using the 'Transform Current Image by Calibration' function under the 'CI-Measurements' tab. This function utilized a user-defined calibration curve to change the coloration of the image itself, therefore calibrating the image to the selected calibration curve.

Another method to ensure that each individual image is calibrated correctly is to generate 'color patches', which are printed strips of varying shades of blue used to determine the calibration of a scanned image. Blue color patches can be created a number of ways, though this study uses Adobe Illustrator to make two rectangles of shades of blue that are visibly lighter and darker from one another. These blue patches are then scanned into CooRecorder along with the IT8.7/2 calibration card. The densitometer function in CooRecorder allows the user to determine the BI value of a user-defined rectangular space. This function, when utilized with a calibrated image of the IT8.7/2 calibration card allows the user to determine the specific BI value of the color patches, which can then be scanned onto each image to ensure the BI values remain consistent for every image (Figure 3.6).


Figure 3.6: Screenshot from CooRecorder displaying blue color patches and densitometer function used for calibration.

Using an EPSON XL12000, cores were scanned with 1200 dpi (dots per inch) resolution and 48-bit color to produce high-definition imagery. The scanned images were individually loaded into CooRecorder 9.2 software for total ring width (TRW) measurement and blue intensity (BI) analysis. CooRecorder delineated growth rings to produce a visually crossdated image. Individual growth rings for each sample were measured to 0.001 mm accuracy. For increment cores from living trees, the incomplete outmost growth ring was created during the year that sampling took place. By assigning a date to the outermost ring, CooRecorder assigned a calendar year to each individual ring. BI data was calculated simultaneously to TRW by recording the reflectance value from a frame structure characterized by user-defined frame specifications. This study used a frame width of 100, a width-limiting factor of 3 , a frame position of 5, a maximum frame deepness of 500, and a relative margin (k) to next ring border. These frame specifications are based on the
specifications utilized by Rydval, et al. (2014).


Figure 3.7: Screenshot from CooRecorder displaying the ring delineation process and BI measurements


Figure 3.8: Screenshot showing the available data outputs for CooRecorder
CooRecorder has several available data outputs for BI measurements. We used the data output methods of raw latewood BI, inverted latewood BI, inverted earlywood BI, and the difference between earlywood and latewood BI (delta BI ). BI values were inverted before exporting so the data could be detrended similarly to detrending methods undergone with maximum latewood density data. All exported time series data was compiled into a collection in CDendro 9.1. CDendro then took the collection of time series data and converted it to Tuscon ring width format to be detrended in ARSTAN. TRW data was crossdated to existing regional chronologies using the software COFECHA to ensure the accuracy of tree ages. In addition to statistically calculating the interseries correlation of the chronology, COFECHA
also records parameters such as average mean sensitivity and flags potential errors in the ring delineation process.

### 3.3.3 Statistical Analysis

The TRW and BI measurements recorded by CooRecorder in this study are processed in the program ARSTAN to detrend the age-related growth trend (Cook and Holmes, 1986) Time series data was then validated using COFECHA, a computer program designed to statistically validate tree ring time series data with existing regional chronologies (Holmes 1983; Grissino-Mayer 2001).

This study utilized a program that performs correlation tests and generates figures to determine the statistical significance of the relationship between tree ring time series data and global climate data. The Royal Netherlands Meteorological Institute (KNMI) developed the web-based application, Climate Explorer, in 1999. To this day, KNMI Climate Explorer is a database of over 10 TB of global climate data and it widely utilized by a number or scientists who work with time series data.

KNMI Climate Explorer takes the time series data that the user inputs and calculates the Pearson Correlation Coefficient between the time series data and the user-selected climate data. KNMI Climate Explorer uses gridded raster data and performs this statistical test repeatedly for every grid cell across a user-defined space. Each Person correlation coefficient value is assigned a color, with positive correlation values being represented with yellows, reds, and oranges, and negative correlation values being represented with blues. These statistical analyses can be generated for every month of the year, or specific months to determine the relationship between time series data and seasonal climate fluctuations.

For this study, we performed our statistical tests using instrumental summer temperature records provided by the Parameter-elevation Relationships on Independent Slopes Model (PRISM) surface temperature dataset, which has a $0.25^{\circ}$ resolution for the contiguous United States and extends from 1895 to 2015.

### 3.4 Results

### 3.4.1 COFECHA

This study contributed to the chronology of PIEN sites that have been sampled in the American Southwest. The locations used for sampling are located on steep gradients and midslope. The oldest PIEN sampled during this study dates back to 1661 , with several other samples dating to the late $17^{\text {th }}$ and early $18^{\text {th }}$ centuries (Figures 3.9 and 3.10). To validate the TRW and BI data used in this study, COFECHA was used to verify our chronology to existing regional chronologies.

The PIEN chronologies developed for this study all express exceptional interseries correlation values. An interseries correlation value of 0.328 is necessary for a $99 \%$ confidence interval (Speer 2010). The TRW time series data presented an interseries correlation of 0.569 . The interseries correlation values for the other two parameter were slightly less, though still significant, with delta BI having a correlation of 0.490 and inverted latewood BI having a correlation of 0.478 .

Based on the range of acceptable mean sensitivity values for climate reconstruction, 0.1-0.4 (Speer 2010), two out of three of the time series datasets used for this study express adequate mean sensitivity, with the exception of inverted latewood BI. TRW had the highest mean sensitivity value at 0.180 , followed by delta BI at 0.129 , and, finally, inverted latewood BI at 0.037.



Figure 3.9: Spaghetti plot displaying trends in TRW for the Wheeler Peak, NM samples

Figure 3.10: Spaghetti plot displaying trends in delta BI for the Wheeler Peak, NM samples

The BI time series data that was generated from CooRecorder were plotted against one another to observe any potential existing relationships between the parameters themselves. A strong positive relationship was observed when comparing inverted latewood BI data and inverted earlywood BI data (Figure 3.11-left). This relationship suggests that greater absorption of blue light in latewood bands of cells results in a greater absorption of blue in earlywood bands of cells. Additionally, a higher density of lignified cells results in a smaller BI value, suggesting the lignin content and cell density is greater with a smaller reflectivity value. A relationship is also present between TRW data and delta BI data (Figure 3.11-right), though it is much weaker than that of inverted latewood BI and inverted earlywood BI. This relationship suggests that BI values are somewhat dependent on annual tree growth. Some outliers are present in both of the scatterplots in Figure 3.11, which may be caused by exceptionally large rings having higher reflectivity and a relatively lower cell density.

The inverted BI parameters and delta BI express similar historical variability (Figure 3.12). Additionally, the start of a growth trend is observable for each of the BI parameters, other than raw values, when nearing the $21^{\text {st }}$ century. When viewing a graph of the PRISM temperature data (Figure 3.13) used in this study, it is evident that this increase in BI follows a global temperature increase. This supports the hypothesis that inverted BI time series data from PIEN at Wheeler Peak, NM is temperature sensitive and follows similar trends as historic temperature variability over the past century.


Figure 3.11: Scatterplots showing the strong positive relationship between inverted earlywood BI and inverted latewood BI (left) and the slightly weaker positive relationship between delta BI and TRW (right)


Figure 3.12: TRW and BI parameter time series 1661-2015 CE.


Figure 3.13: PRISM temperature data centered over the study area showing positive trend over the past several decades.

### 3.4.2 Blue Intensity and Climate Relationships

Our Wheeler Peak, NM tree ring data expressed a statistically significant relationship between the two BI time series used in this study, delta BI and inverted maximum latewood BI, and PRISM instrumental temperature records. To determine seasonal responses in BI parameters, our time series were run against maximum and mean temperature data for August and September separately, averaged maximum and mean temperature data for August through September, and average maximum and mean temperature data for June, July, August, and September (JJAS) temperature data. The figures generated by KNMI Climate Explorer use a blue and red color scheme to display the relationship between the uploaded time series and user-defined climate data, with white showing no correlation, gradually darker blues showing stronger negative correlations, and darker shades of yellows, oranges and reds, showing stronger positive correlations.

### 3.4.2.1 Statistical Analysis

The correlation between the utilized BI parameters and summer temperature data is significant, expressing the most notable positive correlation in the months of August and September. Delta BI and temperature data from the previous year's May, August and September months have a strong positive correlation, with August and September temperature data having a correlation coefficient of greater than 0.25 (Figure 3.32). Correlation coefficients were also generated for current October through January and the previous year's December through March, for forty-year intervals from 1896 to 2015 (Figure 3.33), which also showed that the strongest correlation values were present in the months of August and September for all 82 intervals that were tested. Similar results were found when generating the same figures using the inverted latewood BI parameter, with the previous year's August
and September data have the highest positive correlation with inverted latewood BI time series data (Figures 3.34 and 3.35). When generating these same figures using TRW time series data (Figures 3.36 and 3.37), the signal weakens substantially, with no significant positive relationship present between TRW and any month's temperature data. This supports the hypothesis that BI is a much better predictor of historic temperature variability, as TRW is likely too heavily influenced by other environmental variables that are site-specific, such as moisture availability.


Figure 3.14: Graph showing statistically significant correlations (True) between delta BI parameter and monthly temperature data for current months (Mar-Dec) and the monthly data for the previous year (JAN-OCT).


Figure 3.15: Correlation coefficient values expressed as higher positive correlations as gradually darker blues and higher negative correlations as darker reds. Delta BI time series data ran against current year's monthly temperature data and previous year's monthly temperature data over 40 year intervals from 1896 to 2015.


Figure 3.16: Graph showing statistically significant correlations (True) between inverted latewood BI parameter and monthly temperature data for current months (Mar-Dec) and the monthly data for the previous year (JAN-

OCT).


Figure 3.17: Correlation coefficient values expressed as higher positive correlations as gradually darker blues and higher negative correlations as darker reds. Inverted latewood BI time series data ran against current year's monthly temperature data and previous year's monthly temperature data over 40 year intervals from 1896 to 2015.


Figure 3.18: Graph showing statistically significant correlations (True) between TRW and monthly temperature data for current months (Mar-Dec) and the monthly data for the previous year (JAN-OCT).


Figure 3.19: Correlation coefficient values expressed as higher positive correlations as gradually darker blues and higher negative correlations as darker reds. TRW time series data ran against current year's monthly temperature data and previous year's monthly temperature data over 40 year intervals from 1896 to 2015.

### 3.4.2.1 TRW vs. PRISM Summer Temperature Data

In an effort to emphasize the importance of this temperature proxy at its southern range limit, correlation values between averaged JJAS maximum temperature and TRW (Figure 3.30) and mean temperature and TRW (Figure 3.31) calculated. There is virtually no relationship identified between PRISM mean and maximum JJAS temperature data and TRW, with the exception of some areas in the American Southwest depicting slight spurious negative correlations. Though TRW has been used in existing studies as a proxy for temperature (Esper, 2002), the relationship between the temperature and TRW at Wheeler Peak, NM is nonexistent. It is expected that the lack of temperature signal with TRW is a result of the southern extent of the site, as well as the non-climatic signals produced from disturbances and stand dynamics (Buckley 2018).


Figure 3.20: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM TRW time series and maximum summer (JJAS) temperature


Figure 3.21: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM TRW time series and averaged summer (JJAS) temperature

### 3.4.2.2 Delta BI vs. PRISM Summer Temperature Data

When comparing delta BI to maximum summer temperatures, there is a positive correlation between the two variables, with the strongest temperature response centralized across the American Southwest. When observing the relationship between delta BI and August maximum temperature (Figure 3.14) and delta BI and September maximum temperature (Figure 3.14), the value of the Pearson correlation coefficient is the greatest in areas adjacent to the field site for August, and just north of the field site for September. If a strong correlation value is observed at or near the field site, it signifies that our tree ring time series data is in fact responding to monthly and annual temperature variability. The
temperature signal generated between maximum September (Figure 3.15) temperature and delta BI is not as strong as the relationship with maximum August temperature.

By averaging the maximum annual temperatures for the months of August and September (Figure 3.16), the climate response becomes even stronger as distance from the field site decreases, with the Pearson correlation coefficient between 0.5-0.6 at the field site and extending greater than 0.6 in central and southern Colorado. Though the signal with averaged JJAS (Figure 3.17) maximum temperature is not as prominent, there is still a statistically significant relationship between delta BI and averaged summer maximum temperatures. The temperature response is also not as fixated over the field site, though correlation values of $>0.3$ are observed across much of the American Southwest.


Figure 3.22: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM maximum August temperature


Figure 3.23: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM maximum September temperature


Figure 3.24: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM averaged maximum August-September temperature


Figure 3.25: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM averaged maximum summer (JJAS) temperature

In addition to determining the relationship between maximum temperatures, we also generated results for each tree ring parameter with averaged monthly temperatures. The results from using averaged monthly temperatures were less significant than maximum temperatures, though a temperature response does exist. The output for mean August temperature (Figure 3.18) shows correlation values between 0.2 and 0.3 across New Mexico, southern Colorado, and central Texas, with a few patches of values in excess of 0.3 in northern New Mexico and southern Colorado. Mean September (Figure 3.19) temperatures show a stronger relationship with delta BI at the field site. Most of the areas expressing higher
correlations (0.3-0.4) are located at or near the field site, suggesting that the relationship between delta BI at Wheeler Peak, NM and annual average summer temperatures is positive.

As with delta BI and maximum temperature, when the mean temperatures for August and September (Figure 3.20) are averaged and processed by KNMI Climate Explorer, the temperature signal increases drastically. The signal over the field site and in central Colorado increases to between 0.3 and 0.4 , and the signal for northeastern New Mexico, southeastern Colorado, and Texas increases from between 0.2 and 0.3 to between 0.3 and 0.4. The area comprising correlation values between 0.2 and 0.3 stays relatively consistent, with the aforementioned regions experiencing an increased positive relationship. The mean temperature signal expresses a consistent decline when June, July, August and September (Figure 3.21) temperature data is added, similar to the weakened signal of maximum JJAS averaged temperature. The strong positive temperature signal seen in Figure 3.18 shifts to the East, expressing little to no correlation with the field site, though there is still a positive regional signal across much of the American Southwest.


Figure 3.26: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM mean August temperature


Figure 3.27: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM mean September temperature


Figure 3.28: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM averaged mean August and September temperature


Figure 3.29: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM delta BI time series and PRISM averaged mean summer (JJAS) temperature

### 3.4.2.3 Inverted Latewood BI vs. PRISM Summer Temperature Data

A second BI proxy is utilized in this study to determine if other reflectance values express any significant relationships between temperature and BI. CooRecorder 9.2 offers a number of different data output options, including latewood, earlywood, or full-ring, and raw values, inverted values, and difference between earlywood and latewood BI, or delta BI. We used inverted latewood BI as a second parameter, as maximum latewood blue reflectance values have been shown to produce an annual temperature signal similar to maximum latewood density (Björklund 2014).

The relationship between inverted latewood BI and maximum August (Figure 3.22) temperature is strong, expressing correlation coefficients of 0.2 to 0.5 across many of the Southwestern states. The strongest recorded correlation was observed directly over Wheeler Peak, with a value of 0.4 to 0.5 . A majority of the area of New Mexico and Texas show a statistically significant correlation, and roughly half of Colorado expresses significant correlation.

By averaging August and September (Figure 3.24) temperature data, the positive correlation is shifted back to the west, with a strong correlation of 0.4 to 0.5 at the study site. When observing the relationship, the inverted latewood BI has with averaged maximum JJAS (Figure 3.25) temperature, it is evident that the signal is still present, though it is weaker than the signal for averaged August and September temperature alone.


Figure 3.30: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM maximum August temperature


Figure 3.31: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM maximum September temperature


Figure 3.32: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM averaged maximum August-September temperature


Figure 3.33: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM averaged maximum summer (JJAS) temperature

When observing the relationship between mean PRISM data and inverted latewood BI, it was evident that a signal between the two parameters exists. For August (Figure 3.26) temperature data, the signal is rather weak across the American Southwest, though patches of significant correlation values are still present. A small patch with a correlation value of 0.3 to 0.4 falls directly over the site as well as southern Colorado. The temperature signal for September (Figure 3.27) was much stronger and more widespread, with correlation coefficients of 0.3 to 0.4 directly over the study site as well as central Colorado and parts of Texas. When running inverted latewood BI against averaged August and September (Figure 3.28) mean summer temperature, the positive correlation is strengthened and more centralized
around the study site. At Wheeler Peak, the positive correlation is the highest that is visible, with a correlation coefficient of 0.4 to 0.5 .


Figure 3.34: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM mean August temperature


Figure 3.35: KNMI Climate Explorer output displaying correlations between Wheeler Peak, NM inverted latewood BI time series and PRISM mean September temperature


Figure 3.36: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM averaged mean August and September temperature


Figure 3.37: KNMI Climate Explorer output displaying correlation between Wheeler Peak, NM inverted latewood BI time series and PRISM averaged mean summer (JJAS) temperature

By observing these parameters, we are able to determine which of the tree ring proxies that are used in this study provide the strongest signal for historic temperature. The signal between delta BI and averaged August and September maximum temperatures is much stronger than the temperature signal between TRW and averaged August and September maximum temperature (Figure 3.28). This strong correlation between temperature and delta BI at the field site suggests that BI parameters, specifically delta BI, would be a much more effective predictor of historical temperature variability than TRW.


Figure 3.38: KNMI Climate Explorer outputs comparing correlations between delta BI (left) and TRW (right) parameters against averaged August and September maximum temperature

### 3.5 Discussion

The use of Engelmann spruce for this study allows for the investigation of the variability in historical temperature at Wheeler Peak and the surrounding region. These trees provide exceptionally useful data given that they are located at high latitude and in an alpine environment. Other studies have used different temperature-sensitive tree species, such as Scots pine (Pinus sylvestris L.) (McCarroll et al., 2010; Campbell et al., 2011; Wilson et al.,
2012), Fujian Cypress (Fokienia hodginsii) (Buckley et al., 2018), and bristlecone pine (Pinus aristata) (Salzer et al., 2005), in their reconstructions of historical temperature variability. Previous studies have used networks of tree ring chronologies as well as several different temperature sensitive tree ring parameters including BI, MXD, and TRW (Trouet, et al., 2013; Wilson et al., 2014). These studies had comparable results, finding statistically significant correlations for all of the observed tree ring parameters, though Wilson (2014) found that MXD had a much stronger temperature signal than BI and TRW, which had similar correlations to CRUTS. 3 gridded climate data. However, these results deviate from the finding of our study, as we found that TRW expressed no statistically significant correlation at the field site, whereas BI had strong statistical significance. Additionally, studies using BI to reconstruct temperature in North America experience shortcomings resulting from the relatively short instrumental temperature record in comparison to locations in Europe that have much longer instrumental records.

The methods utilized in this study are comparable to several temperature reconstructions developed over the past several decades. One of the most notable similarities between my results and published studies is the expression of a $20^{\text {th }}$ century warming trend in BI and MXD time series (Salzar et al., 2005; D’Arrigo et al., 2006; McCarroll et al., 2010; Trouet et al., 2013). D’Arrigo et al. (2006) identify the decades that express the warmest reconstructed temperatures, all of which are in the twentieth century in their study as well as four other studies that have been conducted between 1998 and 2005.

In addition to $20^{\text {th }}$ century warming, several existing studies also describe the strongest correlations between BI and MXD time series and gridded temperature data existing during the August and September months (Briffa et al., 2002; Campbell et al., 2011; Buckley et al.,
2018). As the PIEN we sampled express variation in their growth patterns resulting from temperature variability, it is likely that the strongest correlation between BI parameters and instrumental temperature data is during this time due to these months generally experiencing the highest annual temperatures. The studies mentioned previously also use temperaturesensitive tree species in their reconstructions, though PIEN is not the focus species for all of them.

Though Campbell et al. (2011) use slightly different methodologies such as using WinDendro rather than CooRecorder and use a process to remove resinous extractives, they still find that the comparisons between the results derived from MXD and BI data are numerous. The resin removal method utilized by Campbell et al. was experimental, allowing samples to soak in ethanol for different amounts of time before analysis. Samples that have been soaked in ethanol for 30 to 40 hours were found to produce BI results most similar to MXD results, suggesting that, had our samples been soaked in ethanol for this time, our results may have benefited from the samples undergoing the resin removal process. Though resin removal was not utilized in our study, significant correlations between temperature and BI data were still evident.

Wilson et al. $(2014,2016 a)$ suggest that BI studies have a minimum of fourteen series to construct a chronology with an acceptable sample depth. This requirement is much larger than the minimum series depth of eight required by studies using MXD-only chronologies, though the production cost and time required by BI studies are much less. Our study met this threshold and produced results that are comparable to the studies of Wilson et al. (2014, 2016a) Though this minimum requirement of fourteen series has been shown to produce ideal results when reconstructing temperature using BI, several other studies used samples from
twenty or more trees (Babst et al., 2009; Bjorklund et al., 2014; Buckley et al., 2018), suggesting that our study may have benefitted from having a larger sample depth.

Ultimately, we found that there is a clear shift in the growth response of Engelmann spruce at the study site when observing the delta BI dataset. For a majority of the duration of the instrumental temperature record (1895-2003), a positive correlation is present between the delta BI time series and the current year's August and September temperature data. At the turn of the $21^{\text {st }}$ century, this positive correlation with the current year's summer temperatures turns to a negative correlation with the previous year's April-July temperatures, suggesting that the trees at Wheeler Peak, NM are having a negative response to growth stresses occurring in the year prior to ring formation. Additionally, when observing trends both the instrumental temperature data and BI time series, there is a noticeable uptick that occurs near the start of the $21^{\text {st }}$ century, which denotes a warming trend in each dataset. This finding is similar to the growth response to atmospheric warming mentioned in Saladyga and Maxwell's (2015) study looking at climate responses of Eastern hemlock (Tsuga canadensis) in West Virginia, as well as Grissino-Mayer et al.'s (2005) study observing the climate response of ponderosa pine growth.

## CHAPTER 4

## CONCLUSIONS AND FUTURE RESEARCH

Tree ring proxies have been used for the past century to reconstruct historic climatological conditions that extend past instrumental climate data. Tree ring chronologies developed globally have been valid for a number of species that are sensitive to changing climate. Alpine tree species that are common at high latitudes have been shown to express reliable climate signals based on a number of different growth parameters. The different time series data recorded from trees that were sampled for this study was successfully crossdated. This allowed us to validate the identity of the age and associated BI parameters of the Engelmann spruce trees located at Wheeler Peak, NM. The Wheeler Peak chronology extends back to 1661 , with a majority of the sample depth extending back to the 1800 s.

Statistical correlation outputs were derived from each of the BI parameters that were recorded from CooRecorder 9.2. The strongest recorded relationship between a BI parameter and temperature data at the site was with delta BI and averaged August and September maximum temperature. The time series data for each tree ring parameter, delta BI, latewood inverted BI, and TRW, were compared to PRISM temperature data. It was found that a statistically significant relationship exists between maximum and mean summer temperatures and the delta BI and inverted latewood BI parameters at Wheeler Peak, NM as well as regionally across the Midwestern and Southwestern United States.

A warming trend at the turn of the $21^{\text {st }}$ century was present in both the instrumental temperature data as well as the delta BI time series data. This trend was denoted by statistically significant positive correlations for a majority of the duration of the instrumental temperature data followed by the presence of a statistically significant negative correlation
around the year 2003. The trees at the study site were initially producing a positive growth response to the current year's summer temperature, but, due to growth stresses, started expressing a negative growth response to the previous year's summer temperature.

This study has contributed to the consistently growing database of tree ring data for the American Southwest, and has provided useful information on the methods of BI analysis. In the literature, there were potential sources of error that can alter the results of BI analysis, thus, this study can be broadened to test these methods. Rydval (2014) mentions the treatment of samples by soaking them in acetone for varying amount of time can alter the reflectance values by removing any extractives that may be contained in the sample.

To further this study in the future, it would be beneficial to sample from multiple sites in the same region. Two other potential high-elevation PIEN field sites that could strengthen the results of this study are San Leonardo Lakes, NM and Jicarita Peak, NM, both of which are located in the Pecos Wilderness of northern New Mexico. Also, sampling high-elevation sites further south to determine the range that BI studies produce effective results would be valuable. This research adds to the existing global network of BI data, while also targeting an area where no BI data is currently available. By expanding BI research to more locations in the American Southwest where BI data is scarce, there will be further validation of BI methods and the development of laboratories that are able to perform BI analysis.

## REFERENCES

Alexander, R.R., W.D. Shepperd 1990. Picea engelmannii. Pp. 187-203, IN R.M. Burns and B.H. Honkala. Silvics of North America. Volume 1. Conifers. USDA Forest Service Agric. Handbook 654, Washington, D.C.

Anchukaitis, K. J., R. D. D'Arrigo, L. Andreu-Hayles, D. Frank, A. Verstege, A. Curtis, B. M. Buckley, G. C. Jacoby, and E. R. Cook. 2013. Tree-Ring-Reconstructed Summer Temperatures from Northwestern North America during the Last Nine Centuries*. Journal of Climate 26 (10):3001-3012.

Anchukaitis, K., R. Wilson, K. Briffa, U. Büntgen, E. Cook, R. D'Arrigo, N. Davi, J. Esper, D. Frank, B. Gunnarson, G. Hegerl, S. Helama, S. Klesse, P. Krusic, H. Linderholm, V. Myglan, T. Osborn, P. Zhang, M. Rydval, L. Schneider, A. Schurer, G. Wiles, and E. Zorita. 2017. Last millennium Northern Hemisphere summer temperatures from tree rings: Part II, spatially resolved reconstructions. Quaternary Science Reviews 163:1-22.

Babst, F., Frank, D., Buntgen, U., Nievergelt, D., Esper, J., 2009. Effect of sample prepa- ration and scanning resolution on the Blue Reflectance of Picea abies. In: TRACE - Tree Rings in Archeology, Climatology and Ecology, Scientific Technical Report: 09, pp. 189-195.

Baker, W.L.1983. Alpine Vegetation of Wheeler Peak, New Mexico, U.S.A.: Gradient Analysis, Classification, and Biogeography, Arctic and Alpine Research 15 (2):223-240.

Björklund, J. A., B. E. Gunnarson, K. Seftigen, J. Esper, and H. W. Linderholm. 2014. Blue intensity and density from northern Fennoscandian tree rings, exploring the potential to improve summer temperature reconstructions with earlywood information. Climate of the Past 10 (2):877-885.

Briffa, K. R., and F. H. Schweingruber. 1988. Summer temperature patterns over Europe: a reconstruction from 1750 A.D. Based on maximum latewood density indices of conifers. San Diego, CA: Academic Press Inc Elseiver Science.

Briffa, K. R., P. D. Jones, and F. H. Schweingruber. 1992. Tree-Ring Density Reconstructions of Summer Temperature Patterns across Western North America since 1600. Journal of Climate 5 (7):735-754.

Briffa, K.R., Jones, P.D., Schweingruber, F.H. and Osborn, T.J., 1998. Influence of volcanic eruptions on Northern Hemisphere summer temperature over the past 600 years Nature 393 (6684):450.

Briffa, K. R., T. J. Osborn, F. H. Schweingruber, I. C. Harris, P. D. Jones, S. G. Shiyatov, and E. A. Vaganov. 2001. Low-frequency temperature variations from a northern tree ring density network. Journal of Geophysical Research: Atmospheres 106 (D3):2929-2941.

Briffa, K. R., T. J. Osborn, F. H. Schweingruber, P. D. Jones, S. G. Shiyatov, and E. A. Vaganov. 2002. Tree-ring width and density data around the Northern Hemisphere: Part 1, local and
regional climate signals. The Holocene 12 (6):737-757.
Brown, P. M., W. D. Shepperd, C. C. Brown, S. A. Mata, and D. L. Mcclain. 1995. Oldest known Engelmann spruce.

Buckley, B. M., K. G. Hansen, K. L. Griffin, S. Schmiege, R. Oelkers, R. D. D’Arrigo, D. K. Stahle, N. Davi, T. Q. T. Nguyen, C. N. Le, and R. J. Wilson. 2018. Blue intensity from a tropical conifer's annual rings for climate reconstruction: An ecophysiological perspective. Dendrochronologia 50:10-22.

Büntgen, U., D. C. Frank, D. Nievergelt, and J. Esper. 2006. Summer Temperature Variations in the European Alps,a.d.755-2004. Journal of Climate 19 (21):5606-5623.

Campbell, R., D. Mccarroll, N. J. Loader, H. Grudd, I. Robertson, and R. Jalkanen. 2007. Blue intensity in Pinus sylvestris tree-rings: developing a new palaeoclimate proxy. The Holocene 17 (6):821-828

Campbell, R., D. Mccarroll, I. Robertson, N. J. Loader, H. Grudd, and B. Gunnarson. 2011. Blue Intensity In Pinus sylvestris Tree Rings: A Manual for A New Palaeoclimate Proxy. TreeRing Research 67 (2):127-134.

Clark, K. F., Read, C. B., 1972. Geology and ore deposits of Eagle Nest area, New Mexico. New Mexico Bureau of Mines and Mineral Resources Bulletin 94: 152.

Cook, E.R., Holmes, R.L., 1986. Users manual for program ARSTAN. In: Holmes, R.L., Adams, R.K., Fritts, H.C. (Eds.), Tree-Ring Chronologies of Western North America. University of Arizona, Tucson, pp. 50-65.

D'Arrigo, R., R. Wilson, and G. Jacoby. 2006. On the long-term context for late twentieth century warming. Journal of Geophysical Research 111 (D3).

Esper, J. 2002. Low-Frequency Signals in Long Tree-Ring Chronologies for Reconstructing Past Temperature Variability. Science 295 (5563):2250-2253.

Esper, J., F. H. Schweingruber, and M. Winiger. 2002. 1300 years of climatic history for Western Central Asia inferred from tree-rings. The Holocene 12 (3):267-277.

Eschbach, W., P. Nogler, E. Schär, and F. H. Schweingruber, 1995: Technical advances in the radiodensitometrical determination of wood density. Dendrochronologia, 13, 155-168.

Fukazawa, K., 1992. Ultraviolet microscopy. In: Lin, S.Y., Dence, C.W. (Eds.), Methods in Lignin Chemistry. Springer, Berlin, Heidelberg, pp. 110-121.

Grissino-Mayer, H. D. 2001. Evaluating crossdating accuracy: a manual for the program COFECHA. Tree-Ring Research 57, 205-219.

Grissino-Mayer, H. 2003. A manual and tutorial for the proper use of an increment borer. Tree Ring Research, 59 (2), 63-79.

Grissino-Mayer, H., A. Bhuta, M. Crist, J. Doerner, C. Gentry, S. Green, J. Hart, L. Herman, S. Kaplan, R. Keim, E. Larson, D. Mann, W. McCaughey, M. Reddish, S. Stanton, C. Welsh, and D. Wilkins. 2005. Response of Ponderosa Pine to Variable Temporal Scale Environmental Processes, French Creek Drainage, Idaho. 15th Annual North American Dendroecological Fieldweek (NADEF) Final Report.

Holmes, R. L. 1983. Computer assisted quality control in tree-ring dating and measurement. TreeRing Bulletin 43, 69-78.

Hughes, M.K. and Graumlich, L.J., 1996. Multimillennial dendroclimatic studies from the western United States. In Climatic Variations and Forcing Mechanisms of the Last 2000 Years, 109124. Springer, Berlin, Heidelberg.

Knapp, A. K., Smith, W. K. 1982. Factors influencing understory seedling establishment of Engelmann spruce (Picea engelmannii) and subalpine fir (Abies lasiocarpa) in southeast Wyoming. Canadian Journal of Botany 60 (12): 2753-2761.

LaMarche Jr, V.C. and Stockton, C.W., 1974. Chronologies from termperature-sensitive bristlecone pines at upper treeline in Western United States. Tree-Ring Bulletin.

Lange, P.W., 1954. The distribution of lignin in the cell wall of normal and reaction wood from spruce and a few hardwoods. Sven. Papperstidn. 57, 525-532.

Mccarroll, D., E. Pettigrew, A. Luckman, F. Guibal, and J.-L. Edouard. 2002. Blue Reflectance Provides a Surrogate for Latewood Density of High-Latitude Pine Tree Rings. Arctic, Antarctic, and Alpine Research 34 (4):450.

Mccarroll, D., M. Tuovinen, R. Campbell, M. Gagen, H. Grudd, R. Jalkanen, N. J. Loader, and I. Robertson. 2010. A critical evaluation of multi-proxy dendroclimatology in northern Finland. Journal of Quaternary Science 26 (1):7-14.

Menlove, Jim. 2004. Forest Resources of the Carson National Forest. Ogden, UT: U.S. Department of Agriculture, Forest Service, Rocky Mountain Research Station. 16 p.

NCDC. 2018. NOAA Station ID \#USC00298668. Downloaded from National Climatic Data Center, National Oceanic and Atmospheric Administration (http://www.ncdc.noaa.gov/cog/). Last accessed: May 2018.

O'Brien, Renee A. 2003. New Mexico's Forest Resources, 2000. Resour. Bull. RMRS-RB-3. Fort Collins, CO: U.S. Department of Agriculture, Forest Service, Rocky Mountain Research Station. 117 p.

Rydval, M., L. Larsson, L. McGlynn, B.E. Gunnarson, N.J. Loader, G.H.F. Young, and R. Wilson.
2014. Blue intensity for dendroclimatology: Should we have the blues? Experiments from Scotland. Dendrochronologia, 32, 191-204.

Rydval, M., B. E. Gunnarson, N. J. Loader, E. R. Cook, D. L. Druckenbrod, and R. Wilson. 2016. Spatial reconstruction of Scottish summer temperatures from tree rings. International Journal of Climatology 37 (3):1540-1556.

Rydval, M., N. J. Loader, B. E. Gunnarson, D. L. Druckenbrod, H. W. Linderholm, S. G. Moreton, C. V. Wood, and R. Wilson. 2017. Reconstructing 800 years of summer temperatures in Scotland from tree rings. Climate Dynamics 49 (9-10):2951-2974.

Saladyga, T., and R. S. Maxwell. 2015. Temporal Variability in Climate Response of Eastern Hemlock in the Central Appalachian Region. Southeastern Geographer 55 (2):143-163.

Salzer, M.W. and Kipfmueller, K.F., 2005. Reconstructed temperature and precipitation on a millennial timescale from tree-rings in the southern Colorado Plateau, USA. Climatic Change, 70(3):465-487.

Schneider, L., J. E. Smerdon, U. Büntgen, R. J. S. Wilson, V. S. Myglan, A. V. Kirdyanov, and J. Esper. 2015. Revising midlatitude summer temperatures back to A.D. 600 based on a wood density network. Geophysical Research Letters 42 (11):4556-4562.

Schweingruber, F. H., K. R. Briffa, and P. D. Jones. 1991. Yearly maps of summer temperatures in Western-Europe from AD 1750 to AD 1975 and Western-North America from 1600 to 1982results of a results of a radial densitometrical study on tree rings. Vegetatio 92 (1):5-71.

Schweingruber, F. H., and K. R. Briffa. 1996. Tree-Ring Density Networks for Climate Reconstruction. Climatic Variations and Forcing Mechanisms of the Last 2000 Years :43-66.

Sheppard, P.R., Comrie, A.C., Packin, G.D., Angersbach, K. and Hughes, M.K., 2002. The climate of the US Southwest. Climate Research 21 (3):219-238.

Speer, J., 2010. Fundamentals of tree-ring research. University of Arizona Press.
Stoffel, M., M. Khodri, C. Corona, S. Guillet, V. Poulain, S. Bekki, J. Guiot, B. H. Luckman, C. Oppenheimer, N. Lebas, M. Beniston, and V. Masson-Delmotte. 2015. Estimates of volcanicinduced cooling in the Northern Hemisphere over the past 1,500 years. Nature Geoscience 8 (10):784-788.

Stokes, M.A., and T.L. Smiley. 1968. An introduction to tree-ring dating. Chicago, IL: University of Chicago Press.

Stokes, M. A., and T. L. Smiley. 1996. An introduction to tree-ring dating. Tucson, AZ: University of Arizona Press.

Trouet, V., H. F. Diaz, E. R. Wahl, A. E. Viau, R. Graham, N. Graham, and E. R. Cook. 2013. A 1500-year reconstruction of annual mean temperature for temperate North America on decadal-to-multidecadal time scales. Environmental Research Letters 8 (2):024008.

Tucker, J. J. 1979. Estimation of Tree Age Using the Increment Borer. Arboricultural Journal 3 (7):527-531.

Wahl, E.R., Anderson, D.M., Bauer, B.A., Buckner, R., Gille, E.P., Gross, W.S., Hartman, M. and Shah, A., 2010. An archive of high-resolution temperature reconstructions over the past $2+$ millennia. Geochemistry, Geophysics, Geosystems, 11(1).

Wahl, E.R., Diaz, H.F. and Ohlwein, C., 2012. A pollen-based reconstruction of summer temperature in central North America and implications for circulation patterns during medieval times. Global and Planetary Change, 84, pp.66-74.

Walther, G.-R., E. Post, P. Convey, A. Menzel, C. Parmesan, T. J. C. Beebee, J.-M. Fromentin, O. Hoegh-Guldberg, and F. Bairlein. 2002. Ecological responses to recent climate change. Nature 416 (6879):389-395.

Wilson, R. J. S., and B. H. Luckman. 2003. Dendroclimatic reconstruction of maximum summer temperatures from upper treeline sites in Interior British Columbia, Canada. The Holocene 13 (6):851-861.

Wilson, R., D’Arrigo, R, Buckley, B. 2007. A matter of divergence - Tracking recent warming at hemispheric scales using tree-ring data. Journal of Geophysical Research: Atmospheres 112: D17103.

Wilson, R., N. Loader, M. Rydval, H. Patton, A. Frith, C. Mills, A. Crone, C. Edwards, L. Larsson, and B. Gunnarson. 2012. Reconstructing Holocene climate from tree rings: The potential for a long chronology from the Scottish Highlands. The Holocene 22 (1):3-11.

Wilson, R., R. Rao, M. Rydval, C. Wood, L.-Å. Larsson, and B. H. Luckman. 2014. Blue Intensity for dendroclimatology: The BC blues: A case study from British Columbia, Canada. The Holocene 24 (11):1428-1438.

Wilson, R., K. Anchukaitis, K. R. Briffa, U. Büntgen, E. Cook, R. Darrigo, N. Davi, J. Esper, D. Frank, B. Gunnarson, G. Hegerl, S. Helama, S. Klesse, P. J. Krusic, H. W. Linderholm, V. Myglan, T. J. Osborn, M. Rydval, L. Schneider, A. Schurer, G. Wiles, P. Zhang, and E. Zorita. 2016. Last millennium northern hemisphere summer temperatures from tree rings: Part I: The long term context. Quaternary Science Reviews 134:1-18.

Wilson, R., D. Wilson, M. Rydval, A. Crone, U. Büntgen, S. Clark, J. Ehmer, E. Forbes, M. Fuentes, B. E. Gunnarson, H. W. Linderholm, K. Nicolussi, C. Wood, and C. Mills. 2016. Facilitating tree-ring dating of historic conifer timbers using Blue Intensity. Journal of Archaeological Science 78:99-111.

Yanosky, T. M., and C. J. Robinove. 1986. Digital image measurement of the area and anatomical structure of tree rings. Canadian Journal of Botany 64 (12):2896-2902.

Yanosky, T.M., Robinove, C.J., Clark, R.G., 1987. Progress in the image analysis of tree rings. In:

Jacoby, G.C., Hornbeck, J.W. (Eds.), Proceedings, International Symposium on Ecological aspects of Tree-Ring Analysis. National Technical Information Service, Springfield, Virginia, pp. 658-665.

## APPENDIX 1: Wheeler Peak, NM tree ring width COFECHA output

Dendrochronology Program Library Pun ZZ Program COF 17:54 Mon 28 May 2018 Page 1

[^0]QUALITY CONTROL AND DATING CHECK OF TREE-RING MEASUREMENTS
File of DATED series: wle_trw.rwl

CONTENTS:
Part 1: Title page, options selected, summary, absent rings by series Part 2: Histogram of time spans
Part 3: Master series with sample depth and absent rings by year Part 4: Bar plot of Master Dating Series
Part 5: Correlation by segment of each series with Master
Part 6: Potential problems: low correlation, divergent year-to-year changes, absent rings, outliers Part 7: Descriptive statistics

RUN CONTROL OPTIONS SELECTED
VALUE
1 Cubic smoothing spline $50 \%$ wavelength cutoff for filtering 32 years
2 Segments examined are
50 years lagged successively by 25 years
3 Autoregressive model applied
A R
Seri Autoregressive model applied
A Residuals are used in master dating series and testing
4 Series transformed to logarithms
5 CORRELATION is Pearson (parametric, quantitative)
Critical correlation, 99\% confidence level. 3281
6 Master dating series saved
N
7 Ring measurements listed
N
8 Parts printed
1234567
9 Absent rings are omitted from master series and segment correlations (Y)
$\begin{array}{llllll}\text { Time span of Master dating series is } & 1661 \text { to } & 2016 & 356 & \text { years } \\ \text { Continuous time span is } & 1661 \text { to } & 2016 & 356 \text { years }\end{array}$
Portion with two or more series is $\quad 1692$ to $2016 \quad 325$ years
>> WLE10A_I 2005 absent in 1 of 23 series, but is not usually narrow: master index is -.258
*********
C* Number of dated series $23{ }^{*} \mathrm{C}^{*}$
*O* Master series 16612016356 yrs *O*
*F* Total rings in all series 4992 *F*
*E* Total dated rings checked 4961 *E*
*C* Series intercorrelation . 569 *C*
*H* Average mean sensitivity . 180 * $\mathrm{H}^{*}$


|  |  |  | 1707 | -. 929 | 3 | 1757 | . 527 | 6 | 1807 | 2.442 | 11 | 1857 | -. 525 | 18 | 1907 | . 567 | 22 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 1708 | -1.041 | 3 | 1758 | 1.134 | 6 | 1808 | -. 335 | 11 | 1858 | . 330 | 18 | 1908 | -. 531 | 22 |
|  |  |  | 1709 | . 888 | 3 | 1759 | . 475 | 6 | 1809 | . 609 | 11 | 1859 | . 348 | 18 | 1909 | . 180 | 22 |
|  |  |  | 1710 | . 386 | 3 | 1760 | . 147 | 6 | 1810 | -. 132 | 11 | 1860 | -. 459 | 18 | 1910 | -. 830 | 22 |
| 1661 | . 096 | 1 | 1711 | 1.191 | 3 | 1761 | -. 495 | 6 | 1811 | -. 206 | 11 | 1861 | . 154 | 18 | 1911 | . 073 | 22 |
| 1662 | -. 358 | 1 | 1712 | 1.688 | 3 | 1762 | 2.577 | 6 | 1812 | -. 375 | 11 | 1862 | 1.344 | 18 | 1912 | 1.039 | 22 |
| 1663 | 1.709 | 1 | 1713 | 1.448 | 3 | 1763 | . 057 | 6 | 1813 | . 299 | 11 | 1863 | . 552 | 18 | 1913 | . 779 | 22 |
| 1664 | -1.245 | 1 | 1714 | -1.260 | 3 | 1764 | . 466 | 6 | 1814 | . 478 | 11 | 1864 | . 146 | 18 | 1914 | . 079 | 22 |
| 1665 | 2.756 | 1 | 1715 | -2.643 | 3 | 1765 | -. 697 | 6 | 1815 | 1.027 | 12 | 1865 | -. 350 | 18 | 1915 | . 015 | 22 |
| 1666 | . 674 | 1 | 1716 | -1.307 | 3 | 1766 | -. 548 | 6 | 1816 | . 270 | 12 | 1866 | . 248 | 18 | 1916 | . 825 | 22 |
| 1667 | -2.196 | 1 | 1717 | . 835 | 3 | 1767 | . 589 | 6 | 1817 | . 338 | 12 | 1867 | 1.746 | 18 | 1917 | 1.739 | 22 |
| 1668 | -. 574 | 1 | 1718 | . 479 | 3 | 1768 | 2.066 | 6 | 1818 | -1.375 | 12 | 1868 | . 295 | 19 | 1918 | . 526 | 22 |
| 1669 | -1.289 | 1 | 1719 | -. 947 | 3 | 1769 | . 990 | 6 | 1819 | -. 197 | 12 | 1869 | . 742 | 19 | 1919 | -. 368 | 22 |
| 1670 | -2.497 | 1 | 1720 | -. 922 | 3 | 1770 | . 575 | 6 | 1820 | . 208 | 12 | 1870 | -. 432 | 20 | 1920 | -1.867 | 22 |
| 1671 | -. 279 | 1 | 1721 | -1.662 | 3 | 1771 | -. 270 | 6 | 1821 | 1.563 | 14 | 1871 | -. 054 | 20 | 1921 | -1.217 | 22 |
| 1672 | -. 510 | 1 | 1722 | -1.164 | 4 | 1772 | -. 042 | 6 | 1822 | -. 903 | 15 | 1872 | -. 676 | 20 | 1922 | -. 291 | 22 |
| 1673 | . 708 | 1 | 1723 | . 421 | 4 | 1773 | -1.259 | 6 | 1823 | -1.397 | 15 | 1873 | -. 760 | 20 | 1923 | -. 751 | 22 |
| 1674 | . 356 | 1 | 1724 | . 998 | 4 | 1774 | . 045 | 6 | 1824 | -. 453 | 15 | 1874 | 1.256 | 20 | 1924 | 1.106 | 22 |
| 1675 | . 175 | 1 | 1725 | -. 534 | 6 | 1775 | . 409 | 6 | 1825 | -. 413 | 15 | 1875 | . 521 | 20 | 1925 | . 439 | 22 |
| 1676 | -. 225 | 1 | 1726 | . 318 | 6 | 1776 | -1.255 | 6 | 1826 | -. 025 | 15 | 1876 | 1.062 | 20 | 1926 | . 375 | 22 |
| 1677 | -. 099 | 1 | 1727 | . 124 | 6 | 1777 | -1.851 | 6 | 1827 | . 138 | 15 | 1877 | . 014 | 20 | 1927 | . 438 | 22 |
| 1678 | 1.593 | 1 | 1728 | -1.013 | 6 | 1778 | -1.680 | 6 | 1828 | -. 254 | 15 | 1878 | -. 231 | 20 | 1928 | -. 614 | 22 |
| 1679 | 1.474 | 1 | 1729 | -. 254 | 6 | 1779 | -. 633 | 6 | 1829 | . 474 | 16 | 1879 | -1.703 | 20 | 1929 | . 231 | 22 |
| 1680 | . 585 | 1 | 1730 | . 896 | 6 | 1780 | -. 372 | 6 | 1830 | -. 681 | 16 | 1880 | -1.147 | 20 | 1930 | . 017 | 22 |
| 1681 | 2.098 | 1 | 1731 | . 744 | 6 | 1781 | -1.960 | 6 | 1831 | -. 137 | 16 | 1881 | . 477 | 20 | 1931 | -. 138 | 22 |
| 1682 | 1.170 | 1 | 1732 | 1.028 | 6 | 1782 | -. 971 | 6 | 1832 | . 109 | 16 | 1882 | -1.559 | 20 | 1932 | . 668 | 22 |
| 1683 | -. 241 | 1 | 1733 | -. 411 | 6 | 1783 | -. 647 | 6 | 1833 | 1.060 | 16 | 1883 | -1.034 | 20 | 1933 | -. 201 | 22 |
| 1684 | -4.013 | 1 | 1734 | -. 945 | 6 | 1784 | -. 104 | 6 | 1834 | . 313 | 17 | 1884 | -1.180 | 20 | 1934 | -1.068 | 23 |
| 1685 | -4.122 | 1 | 1735 | -. 741 | 6 | 1785 | -. 213 | 7 | 1835 | . 249 | 17 | 1885 | . 560 | 21 | 1935 | -. 125 | 23 |
| 1686 | . 114 | 1 | 1736 | . 574 | 6 | 1786 | . 372 | 7 | 1836 | . 385 | 17 | 1886 | 1.228 | 22 | 1936 | . 264 | 23 |
| 1687 | -. 321 | 1 | 1737 | -1.248 | 6 | 1787 | . 561 | 9 | 1837 | 1.291 | 18 | 1887 | . 370 | 22 | 1937 | -. 575 | 23 |
| 1688 | 1.373 | 1 | 1738 | . 104 | 6 | 1788 | . 891 | 9 | 1838 | . 437 | 18 | 1888 | 1.042 | 22 | 1938 | -. 002 | 23 |
| 1689 | . 863 | 1 | 1739 | . 248 | 6 | 1789 | 1.313 | 9 | 1839 | . 472 | 18 | 1889 | . 952 | 22 | 1939 | -. 343 | 23 |
| 1690 | . 850 | 1 | 1740 | . 401 | 6 | 1790 | . 347 | 9 | 1840 | . 582 | 18 | 1890 | . 443 | 22 | 1940 | -. 494 | 23 |
| 1691 | 2.261 | 1 | 1741 | . 841 | 6 | 1791 | . 551 | 9 | 1841 | -. 630 | 18 | 1891 | . 560 | 22 | 1941 | . 329 | 23 |
| 1692 | . 315 | 2 | 1742 | . 450 | 6 | 1792 | 1.006 | 9 | 1842 | -. 812 | 18 | 1892 | . 703 | 22 | 1942 | . 502 | 23 |
| 1693 | . 234 | 2 | 1743 | -. 134 | 6 | 1793 | . 014 | 9 | 1843 | . 030 | 18 | 1893 | -1.764 | 22 | 1943 | . 913 | 23 |
| 1694 | $-1.281$ | 2 | 1744 | . 860 | 6 | 1794 | . 573 | 9 | 1844 | -. 279 | 18 | 1894 | -. 238 | 22 | 1944 | 1.093 | 23 |
| 1695 | . 020 | 2 | 1745 | 1.944 | 6 | 1795 | -1.089 | 9 | 1845 | -1.555 | 18 | 1895 | . 884 | 22 | 1945 | 1.357 | 23 |
| 1696 | -1.560 | 2 | 1746 | . 118 | 6 | 1796 | -. 342 | 10 | 1846 | -1.737 | 18 | 1896 | -1.144 | 22 | 1946 | -. 435 | 23 |
| 1697 | . 395 | 2 | 1747 | 1.437 | 6 | 1797 | . 008 | 10 | 1847 | -1.011 | 18 | 1897 | -. 119 | 22 | 1947 | 1.318 | 23 |
| 1698 | 2.316 | 2 | 1748 | -2.945 | 6 | 1798 | -. 224 | 11 | 1848 | -. 343 | 18 | 1898 | . 755 | 22 | 1948 | . 666 | 23 |
| 1699 | -. 212 | 3 | 1749 | -. 342 | 6 | 1799 | -. 222 | 11 | 1849 | -. 385 | 18 | 1899 | -1.333 | 22 | 1949 | . 183 | 23 |

PART 3: Master Dating Series:
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| Year | Value | No Ab |  | Year | Value | No | Ab |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1950 | -. 627 | 23 |  | 2000 | . 387 |  | 3 |
| 1951 | 1.443 | 23 |  | 2001 | . 329 |  | 3 |
| 1952 | -2.582 | 23 |  | 2002 | -. 121 |  | 3 |
| 1953 | . 294 | 23 |  | 2003 | -2.763 |  | 3 |
| 1954 | 4.072 | 23 |  | 2004 | -1.810 |  |  |
| 1955 | -. 376 | 23 |  | 2005 | -. 258 | 23 | 1<< |
| 1956 | -1.397 | 23 |  | 2006 - | -1.084 | 23 | 1 |
| 1957 | -2.239 | 23 |  | 2007 | -. 965 |  | 3 |
| 1958 | - . 890 | 23 |  | 2008 | . 729 |  | 3 |
| 1959 | - -.834 | 23 |  | 2009 | 1.353 |  | 3 |
| 1960 | - . 088 | 23 |  | 2010 | . 720 |  | 3 |
| 1961 | -. 098 | 23 |  | 2011 | . 785 |  | 3 |
| 1962 | . 295 | 23 |  | 2012 | . 686 |  |  |
| 1963 | -. 363 | 23 |  | 2013 | 1.245 |  | 3 |
| 1964 | $4-.960$ | - 23 |  | 2014 | -. 184 |  |  |
| 1965 | 5.094 | 23 |  | 2015 | 5.199 |  |  |
| 1966 | 1.323 | 23 |  | 2016 | -1.170 |  |  |
|  |  | 1967 | -. 086 | 8623 |  |  |  |
|  |  | 1968 |  | 5723 |  |  |  |
|  |  | 1969 | 2.419 | 1923 |  |  |  |
|  |  | 1970 | 1.437 | 3723 |  |  |  |
|  |  | 1971 | -. 335 | 335 |  |  |  |
|  |  | 1972 | -. 104 | 0423 |  |  |  |
|  |  | 1973 | . 005 | 0523 |  |  |  |
|  |  | 1974 |  | 5923 |  |  |  |
|  |  | 1975 |  | 5123 |  |  |  |
|  |  | 1976 |  | 0223 |  |  |  |
|  |  | 1977 | -. 427 | 2723 |  |  |  |
|  |  | 1978 | -. 574 | 7423 |  |  |  |
|  |  | 1979 |  | 334 |  |  |  |
|  |  | 1980 |  | 8723 |  |  |  |
|  |  | 1981 | -1.736 | 3623 |  |  |  |
|  |  | 1982 | -. 554 | 5423 |  |  |  |
|  |  | 1983 | -. 607 | 0723 |  |  |  |
|  |  | 1984 | -. 204 | 2423 |  |  |  |
|  |  | 1985 | -. 245 | 4523 |  |  |  |
|  |  | 1986 | -. 492 | 923 |  |  |  |
|  |  | 1987 | -. 993 | 9323 |  |  |  |
|  |  | 1988 | -1.677 | 67723 |  |  |  |
|  |  | 1989 | -. 180 | 8023 |  |  |  |
|  |  | 1990 | . 716 | 1623 |  |  |  |
|  |  | 1991 |  | 50 23 |  |  |  |
|  |  | 1992 | . 234 | 3423 |  |  |  |
|  |  | 1993 | 1.521 | 2123 |  |  |  |




Correlations of 50 -year dated segments, lagged 25 years
Flags: $A=$ correlation under .3281 but highest as dated; $B=$ correlation higher at other than dated position


PART 6: POTENTIAL PROBLEMS:
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For each series with potential problems the following diagnostics may appear:
[A] Correlations with master dating series of flagged 50 -year segments of series filtered with $32-y e a r$ spline,
at every point from ten years earlier ( -10 ) to ten years later (+10) than dated
[B] Effect of those data values which most lower or raise correlation with master series Symbol following year indicates value in series is greater (>) or lesser (<) than master series value
[C] Year-to-year changes very different from the mean change in other series
[D] Absent rings (zero values)
[E] Values which are statistical outliers from mean for the year



|  |  |  | Entire se |  | rrelation ( | 7) |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Lower | 1851> -. 022 | 1891<-. 020 | 1889<-. 019 | $\begin{aligned} & 1927<-.019 \\ & 1821 \text { to } 1870 \end{aligned}$ | $\begin{aligned} & 1837<-.017 \\ & \text { segment: } \end{aligned}$ | 1893> | -. 012 | Higher | 1952 | . 038 | 2003 | . 034 |
| Lower | 1837<-. 079 | 1851> -. 072 | 1866<-. 034 | $\begin{aligned} & 1858<-.033 \\ & 1825 \text { to } 1874 \end{aligned}$ | $\begin{aligned} & \text { 1823> }-.027 \\ & \text { segment: } \end{aligned}$ | 1821< | -. 020 | Higher | 1822 | . 072 | 1850 | . 037 |
| Lower | 1851> -. 087 | 1837<-. 086 | 1866<-. 036 | $\begin{aligned} & 1858<-.035 \\ & 1900 \text { to } 1949 \end{aligned}$ | $\begin{aligned} & 1839<-.013 \\ & \text { segment : } \end{aligned}$ | 1859< | -. 012 | Higher | 1874 | . 043 | 1850 | . 035 |
| Lower | 1927<-. 083 | 1934>-. 042 | 1903<-. 037 | 1945<-. 035 | 1904>-.026 | $1943<$ | -. 026 | Higher | 1902 | . 068 | 1946 | . 032 |

[E] Outliers 23.0 SD above or -4.5 SD below mean for year
$1851+3.7$ SD; $\quad 1934+3.2$ SD



WLE09B 1829 to $2016 \quad 188$ years $\quad$ Series 12


| Lower |  |  | Entire series, effect on correlation ( .528) is: |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 1884<-. 022 | 1910> -. 015 | 1945<-. 014 | 1879> -. 011 | $1840<-.009$ | 1837 < | -. 008 | Higher | 1952 | . 049 | 1893 | . 015 |
|  |  |  |  | 1829 to 1878 | segment: |  |  |  |  |  |  |  |
| Lower | 1840<-. 037 | 1837<-. 032 | 1835<-. 025 | 1842> -. 024 | 1852> -. 018 | 1858< | -. 014 | Higher | 1874 | . 029 | 1853 | . 024 |

[E] Outliers 1 3.0 SD above or -4.5 SD below mean for year
$1910+4.2$ SD




PART 7: DESCRIPTIVE STATISTICS: $\quad 17: 54$ Mon 28 May 2018 Page


| Seq | Series | Inte | val | Years | Segmt | Flags | Master | msmt | msmt | dev | corr | sens | value | dev | corr | () |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | WLE01B_I | 1798 | 2016 | 219 | 9 | 1 | . 482 | 1.05 | 2.22 | . 425 | . 847 | . 178 | 2.60 | . 406 | . 011 | 2 |
| 2 | WLE02A_I | 1834 | 2016 | 183 | 7 | 0 | . 576 | 1.52 | 3.18 | . 651 | . 911 | . 143 | 2.50 | . 327 | -. 012 | 2 |
| 3 | WLE03A | 1886 | 2016 | 131 | 5 | 0 | . 677 | 1.28 | 2.58 | . 525 | . 869 | . 177 | 2.53 | . 342 | . 012 | 1 |
| 4 | WLE04A | 1821 | 2016 | 196 | 8 | 0 | . 668 | 1.69 | 6.52 | 1.025 | . 932 | . 146 | 2.62 | . 386 | . 012 | 1 |
| 5 | WLE04B_I | 1815 | 2016 | 202 | 8 | 3 | . 548 | 1.55 | 3.66 | . 801 | . 926 | . 150 | 2.52 | . 390 | -. 032 | 1 |
| 6 | WLE05A_I | 1787 | 2016 | 230 | 9 | 0 | . 584 | 1.06 | 3.27 | . 573 | . 901 | . 197 | 2.60 | . 370 | . 009 | 1 |
| 7 | WLE05B | 1787 | 2016 | 230 | 9 | 0 | . 615 | 1.04 | 3.70 | . 613 | . 925 | . 186 | 2.63 | . 400 | -. 015 | 2 |
| 8 | WLE07A | 1821 | 2016 | 196 | 8 | 3 | . 527 | 1.86 | 4.32 | . 917 | . 892 | . 159 | 2.74 | . 467 | -. 044 | 1 |
| 9 | WLE07B | 1934 | 2016 | 83 | 3 | 0 | . 528 | 1.80 | 4.32 | . 722 | . 801 | . 204 | 2.69 | . 549 | . 001 | 2 |
| 10 | WLE08A | 1822 | 2016 | 195 | 8 | 1 | . 497 | 1.59 | 3.22 | . 488 | . 691 | . 177 | 2.68 | . 423 | -. 027 | 1 |
| 11 | WLE08B | 1885 | 2016 | 132 | 5 | 0 | . 646 | 1.50 | 3.53 | . 502 | . 772 | . 155 | 2.73 | . 424 | -. 024 | 1 |
| 12 | WLE09B | 1829 | 2016 | 188 | 7 | 1 | . 528 | 1.96 | 6.96 | 1.124 | . 903 | . 156 | 2.53 | . 355 | -. 069 | 1 |
| 13 | WLE10A_I | 1785 | 2016 | 232 | 9 | 1 | . 457 | 1.08 | 2.62 | . 585 | . 912 | . 207 | 2.61 | . 428 | . 022 | 2 |
| 14 | WLE10B | 1870 | 2016 | 147 | 6 | 1 | . 542 | 1.68 | 5.04 | . 772 | . 841 | . 199 | 2.65 | . 333 | -. 015 | 1 |
| 15 | WLE10C | 1837 | 2016 | 180 | 7 | 0 | . 535 | 1.27 | 6.84 | . 973 | . 919 | . 201 | 2.53 | . 303 | -. 013 | 1 |
| 16 | WLE12A | 1868 | 2016 | 149 | 6 | 2 | . 248 | . 58 | 1.45 | . 257 | . 827 | . 222 | 2.73 | . 479 | . 036 | 1 |
| 17 | WLE12B | 1796 | 2016 | 221 | 9 | 0 | . 644 | 1.02 | 2.38 | . 429 | . 913 | . 141 | 2.47 | . 263 | -. 007 | 2 |
| 18 | WLE13A | 1725 | 2016 | 292 | 11 | 0 | . 488 | . 63 | 1.40 | . 248 | . 845 | . 179 | 2.68 | . 482 | -. 066 | 1 |
| 19 | WLE13B | 1722 | 2016 | 295 | 12 | 0 | . 584 | . 73 | 1.51 | . 270 | . 869 | . 153 | 2.66 | . 422 | -. 041 | 1 |
| 20 | WLE15A | 1692 | 2016 | 325 | 13 | 0 | . 606 | . 77 | 1.75 | . 345 | . 861 | . 187 | 2.64 | . 288 | -. 015 | 1 |
| 21 | WLE15B | 1725 | 2016 | 292 | 11 | 0 | . 571 | . 56 | 1.40 | . 245 | . 811 | . 192 | 2.87 | . 453 | -. 014 | 1 |
| 22 | WLE16A | 1699 | 2016 | 318 | 13 | 0 | . 709 | . 75 | 1.82 | . 391 | . 888 | . 204 | 2.50 | . 321 | -. 036 | 1 |
| 23 | WLE16B | 1661 | 2016 | 356 | 13 | 0 | . 650 | . 58 | 1.79 | . 350 | . 874 | . 203 | 2.53 | . 312 | -. 045 | 1 |
| Total or mean: |  |  |  | 499 | 196 |  | 3.569 | 1.10 | 6.96 | . 53 | . 870 | . 180 | 2.87 | . 381 | -. 019 |  |

[^1]
## APPENDIX 2: Wheeler Peak, NM delta blue COFECHA output

PROGRAM COFECHA

QUALITY CONTROL AND DATING CHECK OF TREE-RING MEASUREMENTS
File of DATED series: wle_delta_new.rwl
CONTENTS:
Part 1: Title page, options selected, summary, absent rings by series
Part 2: Histogram of time spans
Part 3: Master series with sample depth and absent rings by year
Part 4: Bar plot of Master Dating Series
Part 5: Correlation by segment of each series with Master
Part 6: Potential problems: low correlation, divergent year-to-year changes, absent rings, outliers
Part 7: Descriptive statistics
CONTROL OPTIONS SELECTED

1 Cubic smoothing spline $50 \%$ wavelength cutoff for filtering
2 Segments examined are
32 years
Autoregressive model applied
50 years lagged successively by 25 years
4 Series transformed to logarithms
A Residuals are used in master dating series and testing
5 CORRELATION is Pearson (parametric, quantitative) Critical correlation, 99\% confidence level . 3281
6 Master dating series saved
N
7 Ring measurements listed
N
8 Parts printed 1234567
9 Absent rings are omitted from master series and segment correlations (Y)
Time span of Master dating series is 1661 to 2015355 years
Continuous time span is 1661 to $2015 \quad 355$ years
Portion with two or more series is 1692 to 2015324 years

${ }^{*} C^{*}$ Number of dated series $24{ }^{*}$ C*
*O* Master series 16612015355 yrs *O*
*F* Total rings in all series 5138 *F*
*E* Total dated rings checked 5107 *E*
*C* Series intercorrelation . 490 *C*
*H* Average mean sensitivity . 129 * $\mathrm{H}^{*}$
*A* Segments, possible problems 53 *A
*** Mean length of series 214.1 ***

ABSENT RINGS listed by SERIES:
(See Master Dating Series for absent rings listed by year)
No ring measurements of zero value




|  |  |  | 1709 | 1.265 | 3 | 1759 | 1.613 | 6 | 1809 | . 490 | 10 | 1859 | . 145 | 19 | 1909 | -. 695 | 23 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 1710 | -. 670 | 3 | 1760 | -1.197 | 6 | 1810 | -. 363 | 10 | 1860 | . 752 | 19 | 1910 | . 809 | 23 |
| 1661 | -. 274 | 1 | 1711 | 1.521 | 3 | 1761 | -1.883 | 6 | 1811 | -. 493 | 10 | 1861 | . 947 | 19 | 1911 | -. 009 | 23 |
| 1662 | -. 463 | 1 | 1712 | . 746 | 3 | 1762 | . 637 | 6 | 1812 | -. 256 | 10 | 1862 | -. 766 | 19 | 1912 | -. 147 | 23 |
| 1663 | -. 813 | 1 | 1713 | 1.215 | 3 | 1763 | . 856 | 6 | 1813 | 1.106 | 10 | 1863 | . 209 | 19 | 1913 | . 817 | 23 |
| 1664 | -. 537 | 1 | 1714 | -1.812 | 3 | 1764 | -. 191 | 6 | 1814 | 1.625 | 10 | 1864 | . 044 | 19 | 1914 | -. 069 | 23 |
| 1665 | -1.573 | 1 | 1715 | -2.772 | 3 | 1765 | . 630 | 6 | 1815 | . 563 | 11 | 1865 | . 212 | 19 | 1915 | -. 306 | 23 |
| 1666 | 3.016 | 1 | 1716 | -. 499 | 3 | 1766 | -. 131 | 6 | 1816 | -. 077 | 11 | 1866 | -1.608 | 19 | 1916 | -. 694 | 23 |
| 1667 | 3.203 | 1 | 1717 | 1.376 | 3 | 1767 | 1.506 | 6 | 1817 | -. 170 | 11 | 1867 | . 329 | 19 | 1917 | . 443 | 23 |
| 1668 | . 166 | 1 | 1718 | . 938 | 3 | 1768 | . 650 | 6 | 1818 | -. 780 | 11 | 1868 | -1.418 | 20 | 1918 | . 005 | 23 |
| 1669 | -. 696 | 1 | 1719 | -. 439 | 3 | 1769 | . 928 | 6 | 1819 | . 182 | 11 | 1869 | . 654 | 20 | 1919 | . 977 | 23 |
| 1670 | -1.345 | 1 | 1720 | -2.429 | 3 | 1770 | 1.115 | 6 | 1820 | . 253 | 12 | 1870 | -. 446 | 22 | 1920 | -2.693 | 23 |
| 1671 | -. 530 | 1 | 1721 | -. 531 | 3 | 1771 | -1.318 | 6 | 1821 | . 499 | 13 | 1871 | . 577 | 22 | 1921 | . 163 | 23 |
| 1672 | -1.007 | 1 | 1722 | -1.545 | 4 | 1772 | -. 791 | 6 | 1822 | . 215 | 14 | 1872 | -. 807 | 22 | 1922 | . 767 | 23 |
| 1673 | . 397 | 1 | 1723 | 1.174 | 4 | 1773 | -. 770 | 6 | 1823 | -. 680 | 15 | 1873 | -. 004 | 22 | 1923 | -2.017 | 23 |
| 1674 | . 467 | 1 | 1724 | . 919 | 4 | 1774 | 1.177 | 6 | 1824 | . 124 | 15 | 1874 | . 550 | 22 | 1924 | 1.379 | 23 |
| 1675 | -. 824 | 1 | 1725 | . 320 | 6 | 1775 | 1.156 | 6 | 1825 | . 385 | 16 | 1875 | 1.049 | 22 | 1925 | . 383 | 23 |
| 1676 | -. 221 | 1 | 1726 | -. 320 | 6 | 1776 | -1.288 | 6 | 1826 | . 070 | 16 | 1876 | . 909 | 22 | 1926 | 1.090 | 23 |
| 1677 | -. 185 | 1 | 1727 | . 614 | 6 | 1777 | -1.456 | 6 | 1827 | -. 478 | 16 | 1877 | 1.157 | 22 | 1927 | . 424 | 23 |
| 1678 | 1.045 | 1 | 1728 | -. 381 | 6 | 1778 | -1.437 | 6 | 1828 | -1.326 | 16 | 1878 | 1.095 | 22 | 1928 | -. 383 | 23 |
| 1679 | -. 100 | 1 | 1729 | 1.395 | 6 | 1779 | -. 548 | 6 | 1829 | -. 185 | 16 | 1879 | -. 041 | 22 | 1929 | -1.091 | 23 |
| 1680 | 1.460 | 1 | 1730 | . 191 | 6 | 1780 | . 210 | 6 | 1830 | . 902 | 17 | 1880 | -. 901 | 22 | 1930 | . 272 | 23 |
| 1681 | . 634 | 1 | 1731 | . 311 | 6 | 1781 | -. 774 | 6 | 1831 | -1.661 | 17 | 1881 | . 299 | 22 | 1931 | . 357 | 23 |
| 1682 | . 677 | 1 | 1732 | -. 654 | 6 | 1782 | -. 368 | 6 | 1832 | -. 482 | 17 | 1882 | -1.219 | 22 | 1932 | . 072 | 23 |
| 1683 | -. 273 | 1 | 1733 | -. 532 | 6 | 1783 | . 307 | 6 | 1833 | . 781 | 17 | 1883 | -. 169 | 22 | 1933 | . 512 | 23 |
| 1684 | -6.727 | 1 | 1734 | -. 737 | 6 | 1784 | . 095 | 6 | 1834 | . 946 | 18 | 1884 | -2.237 | 22 | 1934 | . 395 | 24 |
| 1685 | -3.697 | 1 | 1735 | . 319 | 6 | 1785 | . 190 | 6 | 1835 | . 114 | 18 | 1885 | -1.086 | 23 | 1935 | -1.721 | 24 |
| 1686 | 1.392 | 1 | 1736 | . 180 | 6 | 1786 | -1.041 | 6 | 1836 | -. 072 | 18 | 1886 | -. 605 | 23 | 1936 | . 832 | 24 |
| 1687 | 1.108 | 1 | 1737 | -. 327 | 6 | 1787 | . 817 | 8 | 1837 | . 553 | 19 | 1887 | . 236 | 23 | 1937 | . 430 | 24 |
| 1688 | 1.465 | 1 | 1738 | . 443 | 6 | 1788 | . 348 | 8 | 1838 | -. 508 | 19 | 1888 | -. 058 | 23 | 1938 | -. 889 | 24 |
| 1689 | . 086 | 1 | 1739 | . 148 | 6 | 1789 | -. 365 | 8 | 1839 | -. 314 | 19 | 1889 | 1.211 | 23 | 1939 | . 303 | 24 |
| 1690 | 1.931 | 1 | 1740 | -1.141 | 6 | 1790 | . 787 | 8 | 1840 | . 669 | 19 | 1890 | . 518 | 23 | 1940 | -. 780 | 24 |
| 1691 | . 943 | 1 | 1741 | 1.464 | 6 | 1791 | . 752 | 8 | 1841 | . 200 | 19 | 1891 | . 041 | 23 | 1941 | -1.864 | 24 |
| 1692 | -. 281 | 2 | 1742 | . 354 | 6 | 1792 | . 061 | 8 | 1842 | -. 727 | 19 | 1892 | -. 014 | 23 | 1942 | . 427 | 24 |
| 1693 | 1.085 | 2 | 1743 | -. 715 | 6 | 1793 | . 950 | 8 | 1843 | -. 039 | 19 | 1893 | -. 537 | 23 | 1943 | 1.101 | 24 |
| 1694 | -. 907 | 2 | 1744 | 1.803 | 6 | 1794 | -. 225 | 8 | 1844 | . 477 | 19 | 1894 | -1.155 | 23 | 1944 | . 391 | 24 |
| 1695 | 1.128 | 2 | 1745 | -. 116 | 6 | 1795 | -. 858 | 8 | 1845 | . 138 | 19 | 1895 | . 427 | 23 | 1945 | . 705 | 24 |
| 1696 | -1.532 | 2 | 1746 | -. 402 | 6 | 1796 | -. 766 | 9 | 1846 | -. 117 | 19 | 1896 | . 487 | 23 | 1946 | . 072 | 24 |
| 1697 | . 692 | 2 | 1747 | . 747 | 6 | 1797 | . 387 | 9 | 1847 | -. 428 | 19 | 1897 | . 861 | 23 | 1947 | . 598 | 24 |
| 1698 | . 334 | 2 | 1748 | -1.325 | 6 | 1798 | . 371 | 10 | 1848 | -. 454 | 19 | 1898 | . 670 | 23 | 1948 | . 608 | 24 |
| 1699 | -. 339 | 3 | 1749 | -. 814 | 6 | 1799 | . 267 | 10 | 1849 | -1.252 | 19 | 1899 | . 684 | 23 | 1949 | . 345 | 24 |

PART 3: Master Dating Series:
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| 1950 | -. 625 | 24 | 2000 | . 991 | 24 |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1951 | . 301 | 24 | 2001 | . 548 | 24 |
| 1952 | -. 308 | 24 | 2002 | . 639 | 24 |
| 1953 | 1.066 | 24 | 2003 | -2.136 | 24 |
| 1954 | . 793 | 24 | 2004 | -1.135 | 24 |
| 1955 | -. 206 | 24 | 2005 | -. 125 | 24 |
| 1956 | -. 975 | 24 | 2006 | -1.218 | 24 |
| 1957 | -2.130 | 24 | 2007 | -. 184 | 24 |
| 1958 | . 711 | 24 | 2008 | -1.911 | 24 |
| 1959 | . 090 | 24 | 2009 | . 412 | 24 |
| 1960 | . 880 | 24 | 2010 | . 356 | 24 |
| 1961 | -. 769 | 24 | 2011 | . 789 | 24 |
| 1962 | . 824 | 24 | 2012 | . 941 | 24 |
| 1963 | . 945 | 24 | 2013 | 1.153 | 24 |
| 1964 | . 409 | 24 | 2014 | -. 294 | 24 |
| 1965 | -1.825 | 24 | 2015 | -. 071 | 24 |
| 1966 | . 312 | 24 |  |  |  |
| 1967 | $-1.694$ | 24 |  |  |  |
| 1968 | -1.491 | 24 |  |  |  |
| 1969 | 1.202 | 24 |  |  |  |
| 1970 | . 993 | 24 |  |  |  |
| 1971 | -. 416 | 24 |  |  |  |
| 1972 | . 878 | 24 |  |  |  |
| 1973 | -. 076 | 24 |  |  |  |
| 1974 | . 306 | 24 |  |  |  |
| 1975 | -. 791 | 24 |  |  |  |
| 1976 | . 229 | 24 |  |  |  |
| 1977 | . 873 | 24 |  |  |  |
| 1978 | . 020 | 24 |  |  |  |
| 1979 | -. 086 | 24 |  |  |  |
| 1980 | . 769 | 24 |  |  |  |
| 1981 | . 236 | 24 |  |  |  |
| 1982 | -. 749 | 24 |  |  |  |
| 1983 | . 369 | 24 |  |  |  |
| 1984 | -. 356 | 24 |  |  |  |
| 1985 | . 762 | 24 |  |  |  |
| 1986 | . 001 | 24 |  |  |  |
| 1987 | -. 087 | 24 |  |  |  |
| 1988 | -1.062 | 24 |  |  |  |
| 1989 | . 455 | 24 |  |  |  |
| 1990 | -. 315 | 24 |  |  |  |
| 1991 | -. 647 | 24 |  |  |  |
| 1992 | -. 026 | 24 |  |  |  |
| 1993 | -. 994 | 24 |  |  |  |
| 1994 | 1.056 | 24 |  |  |  |
| 1995 | -1.023 | 24 |  |  |  |


| 1996 | 1.484 | 24 |
| ---: | ---: | ---: |
| 1997 | .128 | 24 |
| 1998 | 1.191 | 24 |
| 1999 | .306 | 24 |



| 1692---a | 1742------A | 1792-----@ | 1842--c | 1892----@ | 1942-------B | 1992----@ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1693---------D | 1743--c | 1793---------D | 1843----@ | 1893--b | 1943---------D | 1993-d |  |
| 1694-d | 1744----------G | 1794----a | 1844-------B | 1894-e | 1944------B | 1994---------D |  |
| 1695---------E | 1745----@ | 1795-c | 1845-----A | 1895-------B | 1945--------C | 1995-d |  |
| 1696 f | 1746---b | 1796--c | 1846----@ | 1896-------B | 1946-----@ | 1996----------F |  |
| 1697--------C | 1747--------C | 1797------B | 1847---b | 1897--------C | 1947-------B | 1997-----A |  |
| 1698------A | 1748-e | 1798------A | 1848---b | 1898--------C | 1948-------B | 1998---------E |  |
| 1699---a | 1749--c | 1799------A | 1849-e | 1899--------C | 1949------A | 1999------A |  |

PARI 5: CORRELATION OF SERIES BY SEGMENTS:
Correlations of 50 -year dated segments, lagged 25 years
Flags: $A=$ correlation under .3281 but highest as dated; $B=$ correlation higher at other than dated position

| Seq | Series Time_span | $\begin{aligned} & 1675 \\ & 1724 \end{aligned}$ | $\begin{aligned} & 1700 \\ & 1749 \end{aligned}$ | $\begin{aligned} & 1725 \\ & 1774 \end{aligned}$ | $\begin{aligned} & 1750 \\ & 1799 \end{aligned}$ | $\begin{aligned} & 1775 \\ & 1824 \end{aligned}$ | 1800 | $\begin{aligned} & 1825 \\ & 1874 \end{aligned}$ | $\begin{aligned} & 1850 \\ & 1899 \end{aligned}$ | $\begin{aligned} & 1875 \\ & 1924 \end{aligned}$ | $\begin{aligned} & 1900 \\ & 1949 \end{aligned}$ | $\begin{aligned} & 1925 \\ & 1974 \end{aligned}$ | $\begin{aligned} & 1950 \\ & 1999 \end{aligned}$ | $\begin{aligned} & 1975 \\ & 2024 \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | WLE01B_I 17982015 |  |  |  |  | . 70 | . 72 | . 51 | . 63 | . 78 | . 61 | . 25 B | . 14 B | . 16 B |
| 2 | WLE02A_I 18342015 |  |  |  |  |  |  | . 10 B | B. 37 | . 85 | . 81 | . 74 | . 75 | . 75 |
| 3 | WLE04A_I 18232015 |  |  |  |  |  | . 22 B | . 22B | B . 34 | . 54 | . 72 | . 67 | . 70 | . 74 |
| 4 | WLE04B_I 18152015 |  |  |  |  |  | . 32 A | . 40 | . 29 A | . 40 | . 50 | . 26 A | . 32 A | . 54 |
| 5 | WLE05A_I 17872015 |  |  |  |  | . 55 | . 64 | . 61 | . 69 | . 60 | . 41 | . 55 | . 57 | . 63 |
| 6 | WLE05B_I 17872015 |  |  |  |  | . 44 | . 60 | . 57 | . 63 | . 42 | . 25 B | . 54 | . 72 | . 62 |
| 7 | WLE07A_N 18212015 |  |  |  |  |  | . 22 A | . 33 A | . 22B | . 45 B | . 61 | . 50 | . 59 | . 72 |
| 8 | WLE07B_N 19342015 |  |  |  |  |  |  |  |  |  |  | . 27 B | . 23 B | . 34 |
| 9 | WLE08A_N 18222015 |  |  |  |  |  | . 18 B | . 22 B | B . 34 | . 68 | . 71 | . 62 | . 55 | . 53 |
| 10 | WLE08B_N 18852015 |  |  |  |  |  |  |  |  | . 61 | . 68 | . 59 | . 60 | . 53 |
| 11 | WLE09B_N 18302015 |  |  |  |  |  |  | . 56 | . 61 | . 62 | . 53 | . 45 B | . 54 | . 53 |
| 12 | WLE10B_N 18702015 |  |  |  |  |  |  |  | . 38 | . 62 | . 69 | . 62 | . 66 | . 54 |
| 13 | WLE10C_N 18372015 |  |  |  |  |  |  | . 13 B | . 23 B | . 60 | . 69 | . 62 | . 62 | . 59 |
| 14 | WLE11A_I 18252015 |  |  |  |  |  |  | . 03B | B. 48 | . 79 | . 77 | . 68 | . 71 | . 76 |
| 15 | WLE11B_N 18202015 |  |  |  |  |  | . 45 | . 45 | . 64 | . 76 | . 76 | . 76 | . 74 | . 75 |
| 16 | WLE11C_N 18702015 |  |  |  |  |  |  |  | . 26 B | . 40 | . 29 A | . 27 A | . 47 | . 41 |
| 17 | WLE12A_N 18682015 |  |  |  |  |  |  |  | -.21B | -.05B | . 22A | . 54 | . 35 | . 22 A |
| 18 | WLE12B_N 17962015 |  |  |  |  | . 62 | . 60 | . 44 | . 63 | . 86 | . 85 | . 72 | . 50 | . 50 |
| 19 | WLE13A_N 17252015 |  |  | .01B | .06B | . 17 B | . 28 A | . 30 A | . 23 B | . 15B | . 19 B | . 54 | . 32 A | . 23 A |
| 20 | WLE13B_N 17222015 |  | . 20 B | . 19 B | . 28 B | . 67 | . 74 | . 68 | . 68 | . 39 | . 30 A | . 49 | . 33 A | . 38 |
| 21 | WLE15A_N 16922015 | . 60 | . 58 | . 48 | . 44 | . 57 | . 75 | . 62 | . 61 | . 59 | . 56 | . 61 | . 49 | . 64 |
| 22 | WLE15B_N 17252015 |  |  | . 44 | . 63 | . 27 B | . 34 B | . 56 | . 65 | . 49 | . 33 | . 60 | . 71 | . 77 |
| 23 | WLE16A_N 16992015 | . 56 | . 53 | . 46 | . 65 | . 24 B | . 07 B | . 22 B | B. 58 | . 66 | . 69 | . 77 | . 67 | . 64 |
| 24 | WLE16B_N 16612015 | . 51 | . 48 | . 36 | . 40 | . 13 B | . 07 B | . 28 B | B. 68 | . 76 | . 80 | . 68 | . 61 | . 68 |
| Av | segment correlation | . 56 | . 45 | . 33 | . 41 | . 44 | . 41 | . 38 | . 45 | . 56 | . 56 | . 55 | . 54 | . 55 |

PART 6: POTENTIAL PROBLEMS:

For each series with potential problems the following diagnostics may appear:
[A] Correlations with master dating series of flagged 50 -year segments of series filtered with $32-y e a r ~ s p l i n e$, at every point from ten years earlier ( -10 ) to ten years later (+10) than dated
[B] Effect of those data values which most lower or raise correlation with master series Symbol following year indicates value in series is greater (>) or lesser (<) than master series value
[C] Year-to-year changes very different from the mean change in other series
[D] Absent rings (zero values)
[E] Values which are statistical outliers from mean for the year

WLEO2A_I 1834 to $2015 \quad 182$ years 2

| [A] | Segment | High | -10 | -9 | -8 | -7 | -6 | -5 | -4 | -3 | -2 | -1 | +0 | +1 | +2 | +3 | +4 | +5 | +6 | +7 | +8 | +9 | +10 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 18341883 | 10 | . 25 | . 05 | . 00 | . 05 | . 03 | 04 | . 06 | . 05 | 11 | 09 | . 10 | . 12 | 07 | 17 | . 00 | 21 | 07 | 25 | . 25 | . 00 | .26* |




[B] Entire series, effect on correlation ( .521) is:

| Lower | 1884> -. 048 | 1836<-. 016 | 1828> -. 011 | 1838> -. 010 | 1854<-.008 | 1835<-.008 | Higher | 1920 | . 017 | 1923 | 017 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1823 to 1872 segment: |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1828> -. 041 | 1836<-. 040 | 1838> -. 033 | 1854<-. 028 | 1858> -. 025 | 1835<-. 022 | Higher | 1866 | . 155 | 1868 | . 035 |
| 1825 to 1874 segment: |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1828> -. 040 | 1836<-. 039 | 1838> -. 033 | 1854<-. 027 | 1858>-. 025 | 1835<-. 022 | Higher | 1866 | . 159 | 1868 | . 037 |

[C] Year-to-year changes diverging by over 4.0 std deviations: 18841885 -4.5 SD
[E] Outliers 23.0 SD above or -4.5 SD below mean for year $1838+3.1 \mathrm{SD} ; \quad 1884+4.8 \mathrm{SD}$

WLE04B_I 1815 to $2015 \quad 201$ years
Series

[C] Year-to-year changes diverging by over 4.0 std deviations: $18841885-4.3$ SD
[E] Outliers 2 3.0 SD above or -4.5 SD below mean for year $1884+5.4$ SD; $\quad 1967+4.2$ SD
WLE05A_I 1787 to $2015 \quad 229$ years
[B] Entire series, effect on correlation ( .572) is: $\begin{array}{llllllllll}\text { Lower } 1920>-.018 & 1990<-.016 & 1834<-.009 & 1931<-.009 & 1938>-.008 \quad 1935>-.007 \text { Higher } 1884 \quad .022 \quad 1805 \quad .020\end{array}$

## WLE05B_I 1787 to $2015 \quad 229$ years


[B] Entire series, effect on correlation ( .474) is:

| Lower | 1791<-. 030 | 1795> -. 010 | 1856<-. 010 | 1915<-. 009 | 1847> -. 009 | 1834<-. 009 | Higher | 1805 | . 051 | 1967 | . 011 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 900 to | 49 segment: |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |

[E] Outliers 43.0 SD above or -4.5 SD below mean for year
$1791-4.6 \mathrm{SD} ; \quad 1847+3.7 \mathrm{SD} ; \quad 1934+3.0 \mathrm{SD} ; \quad 2011+3.8 \mathrm{SD}$

WLE07A_N 1821 to $2015 \quad 195$ years 7

[C] Year-to-year changes diverging by over 4.0 std deviations:
18841885 -4.2 SD
[E] Outliers 1 3.0 SD above or -4.5 SD below mean for year
$1835+3.2$ SD

WLE07B_N 1934 to $2015 \quad 82$ years 8

[B] Entire series, effect on correlation ( .298) is:



| 18701919 | -2 | -. 22 | -. 08 | . 11 | -. 02 | . 11 | . 01 | . 36 | . 06 | . 37 | *-. 07 | . 26 | . 05 | -. 13 | -. 18 | . 01 | -. 04 | -. 08 | -. 17 | . 09 | . 07 | -. 13 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| - - - |  |  | - - - |  | - - | - - | - - | - - | - - - |  | - - - |  | - |  | - - | - - | - - | - - | - - - | - - | - - |  |
| 19001949 | 0 | -. 03 | -. 02 | . 20 | -. 13 | -. 03 | . 05 | -. 14 | . 25 | . 14 | -. 24 | . 29 * | -. 10 | -. 21 | . 09 | -. 14 | -. 08 | -. 04 | -. 03 | . 04 | . 25 | -. 18 |
| 19251974 | 0 | -. 01 | -. 05 | . 21 | -. 06 | . 08 | -. 06 | -. 12 | . 08 | . 15 | . 04 | . $27 *$ | -. 26 | -. 04 | -. 03 | -. 17 | -. 16 | . 06 | -. 03 | . 01 | . 09 | . 09 |

[B] Entire series, effect on correlation (.362) is:

| Lower | 1943<-. 089 | 2003> -. 023 | 1967> -. 015 | 1882> | -. 014 | 1885> | -. 011 | 1909> | -. 009 | Higher | 2008 | . 027 | 1920 | . 025 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1870 to | 1919 segment: |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1882> -. 050 | 1885> -. 040 | 1909>-. 031 | 1903> | -. 017 | 1886< | -. 015 | 1912< | -. 015 | Higher | 1906 | . 031 | 1877 | . 025 |
| 1900 to | 1949 segment: |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1943<-. 262 | 1909>-. 024 | 1940>-. 015 | 1903> | -. 014 | 1912< | -. 012 | 1929> | -. 011 | Higher | 1920 | . 079 | 1923 | . 036 |
| 1925 to | 1974 segment: |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1943<-. 270 | 1967> -. 039 | 1940>-. 015 | 1952> | -. 012 | 1969< | -. 010 | 1929> | -. 010 | Higher | 1957 | . 081 | 1965 | . 035 |

[E] Outliers 1 3.0 SD above or -4.5 SD below mean for year
1943 -5.6 SD


WLE12B_N 1796 to $2015 \quad 220$ years
Series 18
[B] Entire series, effect on correlation ( .626) is:
Lower $1818<-.024 \quad 1866>-.011 \quad 1801>-.010 \quad 1989<-.009 \quad 1851<-.009 \quad 1862>-.008 \quad$ Higher $\quad 1805 \quad .036 \quad 1884 \quad .021$


WLE13B_N 1722 to $2015 \quad 294$ years


[E] Outliers 2 3.0 SD above or -4.5 SD below mean for year
$1727+4.2$ SD; $\quad 1748+3.1$ SD
[B] Entire series, effect on correlation ( .568) is: Lower $1910<-.012 \quad 1701>-.008 \quad 1749>-.007 \quad 1982<-.007 \quad 1771>-.006 \quad 1785<-.006$ Higher $\quad 1805 \quad .018 \quad 1884 \quad .013$
[E] Outliers 2 3.0 SD above or -4.5 SD below mean for year $1695+3.1$ SD; $1783+3.1$ SD

## WLE15B_N 1725 to $2015 \quad 291$ years

Series 22

[B] Entire series, effect on correlation ( .501) is:

| Lower | 1805> -. 025 | 1748> -. 025 | 1742<-. 014 | 1818> -. 007 | 1934<-.007 | 1920> | -. 007 | Higher | 1884 | . 020 | 2003 | . 011 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1775 to 1824 segment: |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | $1805>-.120$ | 1818> -. 040 | 1782<-. 024 | 1785<-. 022 | 1799<-. 022 | 1778> | -. 016 | Higher | 1808 | . 039 | 1814 | . 034 |
| 1800 to | 1849 segment: |  |  |  |  |  |  |  |  |  |  |  |

[E] Outliers 43.0 SD above or -4.5 SD below mean for year



## APPENDIX 3: Wheeler Peak, NM inverted latewood blue intensity COFECHA output

PROGRAM COFECHA

QUALITY CONTROL AND DATING CHECK OF TREE-RING MEASUREMENTS
File of DATED series: wle_lwinv_new.rwl
CONTENTS:
Part 1: Title page, options selected, summary, absent rings by series
Part 2: Histogram of time spans
Part 3: Master series with sample depth and absent rings by year
Part 4: Bar plot of Master Dating Series
Part 5: Correlation by segment of each series with Master
Part 6: Potential problems: low correlation, divergent year-to-year changes, absent rings, outliers
Part 7: Descriptive statistics
CONTROL OPTIONS SELECTED

1 Cubic smoothing spline $50 \%$ wavelength cutoff for filtering
2 Segments examined are
32 years
3 Autoregressive model applied
50 years lagged successively by 25 years
4 Series transformed to logarithms
A Residuals are used in master dating series and testing
4 Series transformed to logarithms Critical correlation, 99\% confidence level .3281
6 Master dating series saved
N
7 Ring measurements listed
N
8 Parts printed 1234567
9 Absent rings are omitted from master series and segment correlations (Y)
Time span of Master dating series is 1661 to $2015 \quad 355$ years
Continuous time span is 1661 to $2015 \quad 355$ years
Portion with two or more series is 1692 to 2015324 years

*C* Number of dated series 24 * $^{*}$
*O* Master series 16612015355 yrs *O*
*F* Total rings in all series 5138 *F*
*E* Total dated rings checked 5107 *E*
*C* Series intercorrelation . 478 * ${ }^{*}$ *
*H* Average mean sensitivity . 037 * $\mathrm{H}^{*}$
*A* Segments, possible problems 50 *A*
*** Mean length of series 214.1 ***

ABSENT RINGS listed by SERIES:
(See Master Dating Series for absent rings listed by year)
No ring measurements of zero value




|  |  |  | 1709 | . 677 | 3 | 1759 | . 101 | 6 | 1809 | . 294 | 10 | 1859 | -. 092 | 19 | 1909 | -1.587 | 23 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  | 1710 | -. 024 | 3 | 1760 | . 148 | 6 | 1810 | -. 970 | 10 | 1860 | . 449 | 19 | 1910 | . 961 | 23 |
| 1661 | -1.028 | 1 | 1711 | 2.068 | 3 | 1761 | $-1.628$ | 6 | 1811 | -. 973 | 10 | 1861 | . 309 | 19 | 1911 | -. 057 | 23 |
| 1662 | . 601 | 1 | 1712 | 1.173 | 3 | 1762 | . 031 | 6 | 1812 | -. 451 | 10 | 1862 | -. 495 | 19 | 1912 | . 191 | 23 |
| 1663 | -1.284 | 1 | 1713 | . 059 | 3 | 1763 | . 335 | 6 | 1813 | 1.061 | 10 | 1863 | . 717 | 19 | 1913 | . 309 | 23 |
| 1664 | -. 871 | 1 | 1714 | -1.619 | 3 | 1764 | -1.358 | 6 | 1814 | 1.584 | 10 | 1864 | -. 050 | 19 | 1914 | -. 079 | 23 |
| 1665 | . 954 | 1 | 1715 | -2.126 | 3 | 1765 | . 114 | 6 | 1815 | -. 071 | 11 | 1865 | . 582 | 19 | 1915 | -. 227 | 23 |
| 1666 | 3.176 | 1 | 1716 | -1.440 | 3 | 1766 | -. 893 | 6 | 1816 | -1.288 | 11 | 1866 | -1.885 | 19 | 1916 | -. 522 | 23 |
| 1667 | 3.386 | 1 | 1717 | . 930 | 3 | 1767 | . 474 | 6 | 1817 | -. 585 | 11 | 1867 | . 170 | 19 | 1917 | . 219 | 23 |
| 1668 | -1.048 | 1 | 1718 | . 922 | 3 | 1768 | . 933 | 6 | 1818 | -. 522 | 11 | 1868 | -1.567 | 20 | 1918 | -. 581 | 23 |
| 1669 | -1.492 | 1 | 1719 | . 299 | 3 | 1769 | 1.556 | 6 | 1819 | . 281 | 11 | 1869 | . 654 | 20 | 1919 | . 910 | 23 |
| 1670 | -1.164 | 1 | 1720 | -. 285 | 3 | 1770 | 1.473 | 6 | 1820 | . 202 | 12 | 1870 | -. 159 | 22 | 1920 | -2.765 | 23 |
| 1671 | -1.134 | 1 | 1721 | -. 300 | 3 | 1771 | -1.720 | 6 | 1821 | . 462 | 13 | 1871 | 1.444 | 22 | 1921 | -. 197 | 23 |
| 1672 | -1.353 | 1 | 1722 | -1.671 | 4 | 1772 | -1.784 | 6 | 1822 | . 230 | 14 | 1872 | -. 327 | 22 | 1922 | . 977 | 23 |
| 1673 | -. 562 | 1 | 1723 | . 800 | 4 | 1773 | -1.124 | 6 | 1823 | . 026 | 15 | 1873 | . 082 | 22 | 1923 | -1.080 | 23 |
| 1674 | -. 360 | 1 | 1724 | -. 509 | 4 | 1774 | . 225 | 6 | 1824 | 1.055 | 15 | 1874 | . 205 | 22 | 1924 | 1.674 | 23 |
| 1675 | -. 592 | 1 | 1725 | -. 238 | 6 | 1775 | 1.087 | 6 | 1825 | 1.440 | 16 | 1875 | . 330 | 22 | 1925 | . 564 | 23 |
| 1676 | 1.070 | 1 | 1726 | -. 507 | 6 | 1776 | -. 578 | 6 | 1826 | . 186 | 16 | 1876 | . 192 | 22 | 1926 | 1.489 | 23 |
| 1677 | . 731 | 1 | 1727 | 1.235 | 6 | 1777 | . 069 | 6 | 1827 | -. 708 | 16 | 1877 | . 088 | 22 | 1927 | . 398 | 23 |
| 1678 | -. 994 | 1 | 1728 | . 929 | 6 | 1778 | -. 358 | 6 | 1828 | -1.182 | 16 | 1878 | . 808 | 22 | 1928 | -. 542 | 23 |
| 1679 | -. 161 | 1 | 1729 | 1.731 | 6 | 1779 | -. 082 | 6 | 1829 | -. 134 | 16 | 1879 | 1.240 | 22 | 1929 | -1.523 | 23 |
| 1680 | 3.307 | 1 | 1730 | -1.147 | 6 | 1780 | 1.750 | 6 | 1830 | . 471 | 17 | 1880 | -. 437 | 22 | 1930 | . 254 | 23 |
| 1681 | 1.726 | 1 | 1731 | -. 766 | 6 | 1781 | -. 277 | 6 | 1831 | -2.318 | 17 | 1881 | . 650 | 22 | 1931 | . 013 | 23 |
| 1682 | 1.010 | 1 | 1732 | -. 548 | 6 | 1782 | $-1.049$ | 6 | 1832 | -. 470 | 17 | 1882 | $-1.335$ | 22 | 1932 | . 366 | 23 |
| 1683 | -1.058 | 1 | 1733 | -. 146 | 6 | 1783 | . 608 | 6 | 1833 | -. 136 | 17 | 1883 | . 793 | 22 | 1933 | . 220 | 23 |
| 1684 | -2.396 | 1 | 1734 | . 460 | 6 | 1784 | -. 133 | 6 | 1834 | 1.363 | 18 | 1884 | -1.957 | 22 | 1934 | . 794 | 24 |
| 1685 | -4.266 | 1 | 1735 | . 873 | 6 | 1785 | . 369 | 6 | 1835 | . 472 | 18 | 1885 | -1.385 | 23 | 1935 | -1.541 | 24 |
| 1686 | -. 204 | 1 | 1736 | 1.196 | 6 | 1786 | -. 712 | 6 | 1836 | -. 039 | 18 | 1886 | -. 922 | 23 | 1936 | . 262 | 24 |
| 1687 | . 085 | 1 | 1737 | -. 020 | 6 | 1787 | 1.496 | 8 | 1837 | . 590 | 19 | 1887 | . 096 | 23 | 1937 | . 138 | 24 |
| 1688 | . 069 | 1 | 1738 | -. 313 | 6 | 1788 | . 237 | 8 | 1838 | -. 820 | 19 | 1888 | -. 458 | 23 | 1938 | -. 776 | 24 |
| 1689 | -1.952 | 1 | 1739 | -. 489 | 6 | 1789 | -. 568 | 8 | 1839 | -. 399 | 19 | 1889 | . 644 | 23 | 1939 | 1.236 | 24 |
| 1690 | 1.920 | 1 | 1740 | $-1.625$ | 6 | 1790 | . 256 | 8 | 1840 | . 019 | 19 | 1890 | . 835 | 23 | 1940 | -. 650 | 24 |
| 1691 | . 605 | 1 | 1741 | . 399 | 6 | 1791 | -. 292 | 8 | 1841 | . 738 | 19 | 1891 | . 033 | 23 | 1941 | -2.065 | 24 |
| 1692 | -. 315 | 2 | 1742 | . 257 | 6 | 1792 | -. 132 | 8 | 1842 | -. 285 | 19 | 1892 | -. 659 | 23 | 1942 | . 352 | 24 |
| 1693 | 1.867 | 2 | 1743 | -. 988 | 6 | 1793 | -. 063 | 8 | 1843 | . 069 | 19 | 1893 | -. 387 | 23 | 1943 | . 534 | 24 |
| 1694 | -. 037 | 2 | 1744 | . 593 | 6 | 1794 | . 832 | 8 | 1844 | -. 142 | 19 | 1894 | -. 253 | 23 | 1944 | . 037 | 24 |
| 1695 | 1.836 | 2 | 1745 | -. 559 | 6 | 1795 | -. 762 | 8 | 1845 | -. 350 | 19 | 1895 | . 037 | 23 | 1945 | . 532 | 24 |
| 1696 | -. 987 | 2 | 1746 | -1.517 | 6 | 1796 | -. 365 | 9 | 1846 | . 310 | 19 | 1896 | . 125 | 23 | 1946 | -. 444 | 24 |
| 1697 | . 816 | 2 | 1747 | . 725 | 6 | 1797 | . 334 | 9 | 1847 | -. 489 | 19 | 1897 | . 878 | 23 | 1947 | . 458 | 24 |
| 1698 | -. 051 | 2 | 1748 | . 293 | 6 | 1798 | 1.303 | 10 | 1848 | . 563 | 19 | 1898 | -. 311 | 23 | 1948 | . 335 | 24 |
| 1699 | -. 404 | 3 | 1749 | 2.089 | 6 | 1799 | . 204 | 10 | 1849 | . 245 | 19 | 1899 | . 868 | 23 | 1949 | -. 075 | 24 |

PART 3: Master Dating Series:
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| 1950 | -. 825 | 24 | 2000 | 1.467 | 24 |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1951 | . 736 | 24 | 2001 | 1.141 | 24 |
| 1952 | . 232 | 24 | 2002 | . 919 | 24 |
| 1953 | 1.190 | 24 | 2003 | -1.671 | 24 |
| 1954 | 1.418 | 24 | 2004 | -. 696 | 24 |
| 1955 | -. 808 | 24 | 2005 | -. 413 | 24 |
| 1956 | -. 623 | 24 | 2006 | -1.820 | 24 |
| 1957 | -2.420 | 24 | 2007 | -. 442 | 24 |
| 1958 | 1.304 | 24 | 2008 | -2.730 | 24 |
| 1959 | . 159 | 24 | 2009 | -. 031 | 24 |
| 1960 | 1.372 | 24 | 2010 | . 195 | 24 |
| 1961 | -. 509 | 24 | 2011 | . 638 | 24 |
| 1962 | . 798 | 24 | 2012 | 1.007 | 24 |
| 1963 | . 944 | 24 | 2013 | 1.313 | 24 |
| 1964 | -. 020 | 24 | 2014 | -. 221 | 24 |
| 1965 | -2.276 | 24 | 2015 | . 802 | 24 |
| 1966 | . 087 | 24 |  |  |  |
| 1967 | -1.964 | 24 |  |  |  |
| 1968 | -1.320 | 24 |  |  |  |
| 1969 | 1.141 | 24 |  |  |  |
| 1970 | 1.034 | 24 |  |  |  |
| 1971 | -. 658 | 24 |  |  |  |
| 1972 | 1.399 | 24 |  |  |  |
| 1973 | -. 180 | 24 |  |  |  |
| 1974 | . 089 | 24 |  |  |  |
| 1975 | -. 776 | 24 |  |  |  |
| 1976 | . 092 | 24 |  |  |  |
| 1977 | . 608 | 24 |  |  |  |
| 1978 | . 096 | 24 |  |  |  |
| 1979 | . 138 | 24 |  |  |  |
| 1980 | 1.018 | 24 |  |  |  |
| 1981 | . 266 | 24 |  |  |  |
| 1982 | -. 605 | 24 |  |  |  |
| 1983 | . 368 | 24 |  |  |  |
| 1984 | -. 260 | 24 |  |  |  |
| 1985 | . 667 | 24 |  |  |  |
| 1986 | . 101 | 24 |  |  |  |
| 1987 | . 147 | 24 |  |  |  |
| 1988 | -1.129 | 24 |  |  |  |
| 1989 | 1.276 | 24 |  |  |  |
| 1990 | -1.192 | 24 |  |  |  |
| 1991 | -1.213 | 24 |  |  |  |
| 1992 | -. 006 | 24 |  |  |  |
| 1993 | -1.042 | 24 |  |  |  |
| 1994 | 1.464 | 24 |  |  |  |
| 1995 | -. 703 | 24 |  |  |  |


| 1996 | 1.367 | 24 |
| ---: | ---: | ---: |
| 1997 | .038 | 24 |
| 1998 | .948 | 24 |
| 1999 | -.037 | 24 |



| 1691--------B | 1741------ ${ }^{\text {B }}$ | 1791---a | 1841--------C | 1891-----@ |
| :---: | :---: | :---: | :---: | :---: |
| 1692---a | 1742------A | 1792----a | 1842---a | 1892--c |
| 1693----------G | 1743-d | 1793----@ | 1843-----@ | 1893---b |
| 1694----@ | 1744--------B | 1794--------C | 1844----a | 1894----a |
| 1695----------G | 1745--b | 1795--c | 1845---a | 1895-----@ |
| 1696-d | 1746 f | 1796---a | 1846------A | 1896-----@ |
| 1697--------C | 1747--------C | 1797-------A | 1847---b | 1897--------D |
| 1698----@ | 1748------A | 1798----------E | 1848-------B | 1898---a |
| 1699---b | 1749----------H | 1799------A | 1849------A | 1899--------C |




PART 5: CORRELATION OF SERIES BY SEGMENTS:
Correlations of 50 -year dated segments, lagged 25 years
Flags: $A=$ correlation under .3281 but highest as dated; $B=$ correlation higher at other than dated position
$\begin{array}{lllllllllllllllllll}\text { Seq Series } & \text { Time_span } & 1675 & 1700 & 1725 & 1750 & 1775 & 1800 & 1825 & 1850 & 1875 & 1900 & 1925 & 1950 & 1975 \\ & & 1724 & 1749 & 1774 & 1799 & 1824 & 1849 & 1874 & 1899 & 1924 & 1949 & 1974 & 1999 & 2024\end{array}$

WTE01B I 17982015
2 WIEO2A I 18342015
2 WLE02A_I 18342015
WLE04A_I 18232015
4 WLE04B_I 18152015
6 WLE05B_I 17872015
7 WLE07A_N 18212015
8 WLE07B_N 19342015
9 WLE08A_N 18222015
10 WLE08B_N 18852015
11 WLE09B_N 18302015 12 WLE10B N 18702015 13 WLE10C_N 18372015 14 WLE11A_I 18252015 15 WLE11B_N 18202015 16 WLE11C_N 18702015 17 WLE12A N 18682015 18 WLE12B_N 17962015 19 WLE13A N 17252015 20 WLE13B N 17222015 $\begin{array}{llll}20 & \text { WLE13B_N } 1722 & 2015 \\ 21 \text { WLE15A_N } 1692 & 2015\end{array}$ 22 WLE15B_N 17252015 23 WLE16A_N 16992015 24 WLE16B_N 16612015 Av segment correlation

|  |  |  |  | . 51 | . 54 | . 60 | . 66 | . 76 | . 57 | . 33A | . 32 A | . 27A |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  | . 44 | . 58 | . 87 | . 67 | . 56 | . 69 | . 69 |
|  |  |  |  |  | . 42 | . 46 | . 45 | . 46 | . 67 | . 69 | . 72 | . 69 |
|  |  |  |  |  | . 37 | . 40 | . 27 A | . 38 | . 49 | . 42 | . 50 | . 64 |
|  |  |  |  | . 64 | . 63 | . 65 | . 63 | . 56 | . 43 | . 59 | . 66 | . 63 |
|  |  |  |  | . 53 | . 58 | . 61 | . 76 | . 54 | . 27A | . 59 | . 70 | . 58 |
|  |  |  |  |  | . 33 | . 44 | . 40 | . 52 | . 64 | . 67 | . 73 | . 74 |
|  |  |  |  |  |  |  |  |  |  | . 31B | . 35 | . 46 |
|  |  |  |  |  | . 21B | . 21B | . 22B | . 57 | . 67 | . 66 | . 54 | . 57 |
|  |  |  |  |  |  |  |  | . 60 | . 66 | . 62 | . 61 | . 57 |
|  |  |  |  |  |  | . 56 | . 38 | . 39 | . 49 | . 42 B | . 55 | . 61 |
|  |  |  |  |  |  |  | -.05B | . 20B | . 60 | . 56 | . 58 | . 61 |
|  |  |  |  |  |  | -. 05B | -.01B | . 33 | . 53 | . 56 | . 65 | . 70 |
|  |  |  |  |  |  | -.16B | . 29B | . 59 | . 60 | . 62 | . 71 | . 77 |
|  |  |  |  |  | . 02B | . 03B | . 45 | . 75 | . 72 | . 72 | . 75 | . 79 |
|  |  |  |  |  |  |  | . 24 B | . 39 | . 28 A | . 34 | . 49 | . 44 |
|  |  |  |  |  |  |  | .01B | . 22B | . 50 | . 51 | . 32 A | . 20 A |
|  |  |  |  | . 53 | . 54 | . 50 | . 48 | . 71 | . 76 | . 65 | . 37 | . 41 |
|  |  | . 23 B | . 20B | . 04 B | . 39 | . 43 | . 21B | . 02B | . 10B | . 57 | . 37 | . 33 A |
|  | . 39 | . 42 | . 41 | . 36 | . 63 | . 59 | . 63 | . 37 | . 20B | . 60 | . 50 | . 50 |
| . 20B | . 24 B | . 26 B | . 22A | . 42 | . 64 | . 57 | . 63 | . 62 | . 46 | . 63 | . 62 | . 59 |
|  |  | . 44 | . 23B | . 22A | . 45 | . 48 | . 67 | . 53 | . 31 A | . 60 | . 60 | . 57 |
| . 29A | . 33 | . 50 | . 56 | . 23B | .17B | . 30 A | . 47 | . 35 | . 44 | . 74 | . 81 | . 81 |
| . 16 B | . 19B | . 33 A | . 29 B | . 18B | . 08B | . 32 B | . 72 | . 77 | . 70 | . 69 | . 70 | . 77 |
| . 22 | . 29 | . 36 | . 32 | . 37 | . 40 | . 39 | . 41 | . 50 | . 51 | . 57 | . 58 | . 58 |

For each series with potential problems the following diagnostics may appear:
[A] Correlations with master dating series of flagged 50 -year segments of series filtered with $32-y e a r$ spline, at every point from ten years earlier ( -10 ) to ten years later ( +10 ) than dated
[B] Effect of those data values which most lower or raise correlation with master series Symbol following year indicates value in series is greater (>) or lesser (<) than master series value
[C] Year-to-year changes very different from the mean change in other series
[D] Absent rings (zero values)
[E] Values which are statistical outliers from mean for the year


[E] Outliers $1 \quad 3.0 \mathrm{SD}$ above or -4.5 SD below mean for year $1884+4.5$ SD


WLE05B_I 1787 to $2015 \quad 229$ years 6

| [A] | Segment | High | -10 | -9 | -8 | -7 | -6 | -5 | -4 | -3 | -2 | -1 | +0 | +1 | +2 | +3 | +4 | +5 | +6 | +7 | +8 | +9 | +10 |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 19001949 | 0 | . 14 | . 05 | . 24 | . 11 | -. 14 | -. 03 | -. 25 | -. 23 | . 18 | . 06 | .27* | . 06 | -. 08 | -. 02 | -. 09 | . 18 | . 01 | -. 04 | . $00-$ | -. 16 | -. 06 |  |
| [B] | Entire series, effect <br> Lower 1792> -. 013 |  |  | on correlation$1847>-.012$ |  |  | ( .523) is: |  |  | 1930< | -. 01 |  | 1828< | -. 010 |  | 1920> | -. 009 | Hig |  | 1884 | . 024 |  | 1831 | . 014 |
|  | 1900 to Lower | 949 se | nent: |  | $3>$ | 032 |  | 5> - | 032 | 1915< | -. 02 |  | 1938> | -. 021 |  | 1943< | -. 012 | Hig | er | 1929 | . 034 |  | 1939 | . 025 |

[E] Outliers $3 \quad 3.0 \mathrm{SD}$ above or -4.5 SD below mean for year $1792+3.0$ SD; $\quad 1847+4.0$ SD; $\quad 2011+4.0$ SD

[E] Outliers 1 3.0 SD above or -4.5 SD below mean for year
$1835+3.1$ SD



WLE11A_I 1825 to $2015 \quad 191$ years $\quad$ Series 14

[B] Entire series, effect on correlation ( .486) is:


## WLE11B_N 1820 to $2015 \quad 196$ years

Series 15

[B] Entire series, effect on correlation ( .579) is:

| Lower | 1828> -. 028 | 1850> -. 012 | 1866> -. 010 | 1837<-.009 | 1855<-. 009 | 2013<-. 007 | Higher | 1884 | . 024 | 1920 | . 023 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1820 to | 1869 segment: |  |  |  |  |  |  |  |  |  |  |
| Lower | 1828> -. 094 | 1850> -. 037 | 1837<-.028 | 1855<-. 023 | 1861<-. 019 | 1822<-. 018 | Higher | 1831 | . 060 | 1868 | . 050 |
| 1825 to | 1874 segment: |  |  |  |  |  |  |  |  |  |  |
| Lower | 1828> -. 093 | 1850>-. 036 | 1837<-. 028 | 1855<-. 023 | 1861<-. 020 | 1862>-. 015 | Higher | 1831 | . 056 | 1868 | . 047 |

[E] Outliers $3 \quad 3.0 \mathrm{SD}$ above or -4.5 SD below mean for year
$1828+3.7 \mathrm{SD} ; \quad 1849+3.0 \mathrm{SD} ; \quad 1850+3.5 \mathrm{SD}$
===========================================================================================================================================1
WLE11C_N 1870 to $2015 \quad 146$ years


[B] Entire series, effect on correlation ( .386) is:

| Lower | 1943<-. 085 | 2003> -. 017 | 1882> -. 015 | 1967> | -. 013 | 1885> | -. 009 | 1990> | -. 009 | Higher | 2008 | . 058 | 1957 | . 034 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1870 to 1919 segment: |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1882> -. 057 | 1885> -. 035 | 1918> -. 032 | 1903> | -. 023 | 1912< | -. 014 | 1902< | -. 013 | Higher | 1910 | . 029 | 1909 | . 026 |
| 1900 to | 1949 segment: |  |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1943<-. 247 | 1918> -. 019 | 1940> -. 015 | 1903> | -. 015 | 1927< | -. 009 | 1912< | -. 008 | Higher | 1920 | . 057 | 1929 | . 040 |

[E] Outliers 23.0 SD above or -4.5 SD below mean for year
1918 +3.0 SD; 1943 -5.1 SD

WLE12A_N 1868 to $2015 \quad 148$ years
Series 17


WLE13A_N 1725 to 2015 years 291 Series 19

|  | Segment | High | -10 | -9 | -8 | -7 | -6 | -5 | -4 | -3 | -2 | -1 | +0 | +1 | +2 | +3 | +4 | +5 | +6 | +7 | +8 | +9 | +10 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | 17251774 | -3 | . 21 | -. 05 | -. 04 | -. 16 | . 24 | . 09 | -. 18 | . $30 *$ | -. 11 | -. 08 | . 23 | . 07 | . 07 | -. 15 | -. 18 | . 00 | -. 19 | . 12 | . 19 | -. 20 | -. 08 |
|  | 17501799 | 7 | . 20 | -. 12 | -. 15 | . 03 | . 03 | . 07 | -. 05 | . 14 | -. 13 | -. 10 | . 20 | . 06 | . 06 | -. 16 | -. 15 | -. 01 | -. 15 | . 26 * | . 11 | -. 09 | . 23 |
|  | 17751824 | 10 | . 21 | -. 13 | -. 21 | . 11 | -. 04 | . 07 | . 13 | . 01 | -. 22 | . 06 | . 04 | -. 13 | -. 08 | -. 17 | -. 13 | . 16 | . 15 | . 25 | -. 04 | -. 02 | . $36 *$ |
|  | - - - - - |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
|  | 18501899 | 2 | . 03 | -. 14 | -. 09 | -. 11 | . 00 | -. 25 | -. 03 | -. 31 | . 27 | -. 25 | . 21 | -. 11 | . 28 * | -. 02 | . 25 | . 12 | . 26 | . 11 | . 19 | . 11 | . 07 |
|  | 18751924 | 3 | . 00 | -. 21 | -. 03 | -. 24 | -. 08 | -. 23 | -. 02 | -. 46 | . 13 | . 00 | . 02 | . 06 | . 17 | . 21 * | . 17 | -. 03 | . 10 | . 05 | -. 01 | . 10 | . 05 |
|  | 19001949 | -8 | . 05 | -. 20 | . $18 *$ | . 15 | . 04 | . 16 | -. 28 | -. 31 | . 10 | . 07 | . 10 | . 09 | -. 02 | . 14 | . 07 | -. 14 | . 12 | -. 05 | -. 21 | . 05 | -. 10 |
|  |  | - - - | - - | - - | - - | - - | - - | - - | - | - | - | - - | - - | - - | - - | - - - | - - | - - | - - | - - | - - | - - | - - |
|  |  | 0 | . 07 | -. 05 | -. 16 | -. 29 | . 16 | -. 10 | -. 07 | . 09 | . 10 | -. 14 | . $33 *$ | - | - | - | - | - | - | - | - | - |  |

[B] Entire series, effect on correlation ( .260) is:

| Lower 1923> -. 020 | 1990>-. 010 | 1901<-. 009 | 1860<-. 009 | 1798<-. 009 | 1737> -. 007 | Higher | 1828 | . 015 | 1957 | . 014 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1725 to 1774 segment: |  |  |  |  |  |  |  |  |  |  |
| Lower 1737> -. 040 | 1730> -. 034 | 1742<-.031 | 1773> -. 027 | 1748> -. 018 | 1760<-. 017 | Higher | 1771 | . 053 | 1740 | . 031 |
| 1750 to 1799 segment: |  |  |  |  |  |  |  |  |  |  |
| Lower 1798<-.041 | 1773>-.028 | 1783<-. 024 | 1779>-. 016 | 1760<-. 015 | 1769<-.013 | Higher | 1771 | . 055 | 1764 | . 039 |

1798<-.041
1775 to 1824 segment:

| Lower | 1798<-. 052 | 1809<-. 036 | 1817> -. 033 | 1783<-. 032 | 1779> | -. 016 | 1800> | -. 015 | Higher | 1807 | . 041 | 1795 | . 040 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1850 to | 1899 segment: |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | $1860<-.077$ | 1879<-. 041 | 1881<-. 035 | 1885> -. 032 | 1850> | -. 029 | 1875< | -. 016 | Higher | 1884 | . 067 | 1866 | . 059 |
| 1875 to | 1924 segment: |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1923> -. 104 | 1901<-. 047 | 1885> -. 027 | 1879<-. 023 | 1905< | -. 021 | 1881< | -. 021 | Higher | 1884 | . 069 | 1920 | . 061 |
| 1900 to | 1949 segment: |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1923>-. 117 | 1901<-. 048 | 1905<-. 022 | 1929>-. 015 | 1917< | -. 014 | 1943< | -. 013 | Higher | 1935 | . 067 | 1920 | . 045 |
| 1966 to | 2015 segment: |  |  |  |  |  |  |  |  |  |  |  |  |
| Lower | 1990> -. 067 | 1968> -. 046 | 1975> -. 033 | 1996<-. 026 | 1993> | -. 022 | 1985< | -. 021 | Higher | 1967 | . 078 | 2006 | . 063 |

[C] Year-to-year changes diverging by over 4.0 std deviations:
19221923 4.4 SD 19231924 -5.1 SD
[E] Outliers $4 \quad 3.0 \mathrm{SD}$ above or -4.5 SD below mean for year
$1752+3.2 \mathrm{SD} ; \quad 1817+3.9 \mathrm{SD} ; \quad 1923+6.6 \mathrm{SD} ; \quad 1990+3.2 \mathrm{SD}$

WLE13B_N 1722 to 2015 years 294

[B] Entire series, effect on correlation ( .454) is:
Lower $1902<-.0321965>-.011 \quad 1748>-.011 \quad 1737>-.010 \quad 1993>-.008 \quad 1903>-.006$ Higher $\quad 1884 \quad .027 \quad 1831 \quad .019$

| 1900 to 1949 segment: |  |
| ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| Lower $1902<-.181$ | $1903>-.030 \quad 1901<-.026 \quad 1922<-.024 \quad 1929>-.019 \quad 1911<-.013$ Higher $1920 \quad .083 \quad 1924 \quad .038$ |

[C] Year-to-year changes diverging by over 4.0 std deviations:
19021903 5.0 SD
[E] Outliers 3 3.0 SD above or -4.5 SD below mean for year
$1748+3.8 \mathrm{SD} ; \quad 1902-4.8 \mathrm{SD} ; \quad 1993+3.3 \mathrm{SD}$

## WLE15A_N 1692 to 2015324 years



 $\begin{array}{rllrrrrrrrrrrrrrrrrrrrrrrrrrrr}1725 & 1774 & 6 & -.07 & -.29 & .06 & .30 & -.03 & -.04 & -.12 & -.31 & -.13 & -.04 & .26 & .02 & -.11 & .27 & -.20 & -.03 & .34 *-.15 & .10 & .01 & .03 \\ 1750 & 1799 & 0 & .07 & -.25 & .09 & .14 & .02 & .19 & -.12 & -.19 & -.13 & .10 & .22 *-.11 & -.12 & .08 & -.04 & .04 & .18 & -.16 & -.14 & -.21 & .08\end{array}$
[B] Entire series, effect on correlation ( .457) is:

| Lower 1737< -. 020 | 1716> -. 011 | 1929> -. 009 | 1771> -. 008 | 1764> -. 008 | 1698> | -. 008 | Higher | 1884 | . 025 | 1957 | . 017 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1692 to 1741 segment: |  |  |  |  |  |  |  |  |  |  |  |
| Lower 1737<-. 082 | 1716> -. 039 | 1698>-. 037 | 1694<-.033 | 1713<-. 032 | 1721< | -. 020 | Higher | 1730 | . 034 | 1722 | . 028 |
| 1700 to 1749 segment: |  |  |  |  |  |  |  |  |  |  |  |
| Lower 1737<-.109 | 1713<-. 039 | 1716> -. 038 | 1743> -. 034 | 1721<-. 026 | 1700< | -. 018 | Higher | 1749 | . 033 | 1730 | . 032 |

1725 to 1774 segment:
Lower $1737<-.1091743>-.0391764>-.038$ 1771>-.038 1750<-.025 1752<-.017 Higher $1761 \quad .043 \quad 1749.035$
1750 to 1799 segment:
Lower 1792> - 0
$1764>-.0431771>-.041 \quad 1794<-.039 \quad 1750<-.031 \quad 1752<-.021$ Higher $1761 \quad .061 \quad 1758$. 045

[E] Outliers 1 3.0 SD above or -4.5 SD below mean for year $1805+3.2$ SD
[*] Early part of series cannot be checked from 1661 to 1691 -- not matched by another series


[C] Year-to-year changes diverging by over 4.0 std deviations:
16941695 -4.5 SD
[E] Outliers 3 . 3.0 SD above or -4.5 SD below mean for year
$1694+3.0$ SD; $\quad 1759+3.0$ SD; $\quad 1831+3.7$ SD



[^0]:    PROGRAM COFECHA

[^1]:    - = [ COFECHA ZZ COF ] = -

