Does Bedrock Strength Control Valley Morphometry? Examples from Big Creek, Central Idaho.

> Zachery M. Lifton 4-6-05

# **Hypothesis**

Bedrock with low strength will produce wide valley floors, low river gradients, and low hillslope gradients. Bedrock with high strength will produce narrow valley floors, steep river gradients, and steep hillslope gradients.

# Approach

In this study, a quantitative measure of bedrock strength is compared to three valley parameters: 1) Valley Floor Width 2) Stream Gradient 3) Hillslope Gradient

### Approach

Intuitively, rock type must influence the development of river valleys. "Soft" or "weak" rock erodes more easily than "hard" or "strong" rock.

But what do "soft", "hard", "strong", or "weak" mean?



### Approach

Additionally, rock type or lithology alone does not provide enough information for classifying, explaining, or predicting valley morphometry.



Something else is needed...







Reach selection: Uniformly wide sections Generally avoid major tributaries Range in length from ~200m – 2500m



Rock strength measured with a Schmidt Hammer.
Valley floor width measured with a laser rangefinder.

 Stream gradient measured from DEM in ArcMap and RiverTools.
 Hillslope gradient measured from DEM in

ArcMap.

The Schmidt Hammer "measures the distance of rebound of a controlled impact on a rock surface" (Day, 1980).





"Elastic recovery depends on surface hardness, and hardness is related to mechanical strength" (Day, 1980), therefore rebound measures relative hardness and may be thought of as a proxy for overall mechanical strength.

Furthermore, "surface hardness [...] may be a better measure of resistance to erosion than the bulk compressive strength" (Day, 1980).



Variance analysis shows that the mean of rebound readings at each reach are distinct: the variance between reaches is higher than the variance within the reaches.





P = 7.859 E-05



P = 6.612 E-06



P = 0.0547



P = 0.0212



P = 1.082 E-05



The relationship between relative rock strength (as measured by the Schmidt Hammer) and valley floor width is described by a power law function. **Sklar and Dietrich** (2001) found a similar relationship in an abrasion study.

There also appears to be an aspect effect on rock strength: the south side of the valley (north facing slopes) generally has higher rock strength.



#### t-test of all North data vs. all South data

#### N/S compared by ALL DATA

t-Test: Two-Sample Assuming Unequal Variances

|                              | All North | All South |  |
|------------------------------|-----------|-----------|--|
| Mean                         | 1.620513  | 1.654133  |  |
| Variance                     | 0.029982  | 0.015141  |  |
| Observations                 | 1078      | 1231      |  |
| Hypothesized Mean Difference | 0         |           |  |
| df                           | 1913      |           |  |
| t Stat                       | -5.30824  |           |  |
| P(T<=t) one-tail             | 6.18E-08  |           |  |
| t Critical one-tail          | 1.645651  |           |  |
| P(T<=t) two-tail             | 1.24E-07  |           |  |
| t Critical two-tail          | 1.961207  |           |  |











Local stream gradient doesn't seem to be controlled by rock strength, probably because Big Creek is an alluvial stream.

#### Conclusions

Rock strength is related to valley floor width by a power-law relationship. Stream gradient does not seem to be affected by rock strength (in an alluvial river). Rock strength is influenced by aspect. North-facing slopes have higher strength than south-facing slopes. Melton (1960) offers some explanations, but this study did not address that question directly.

Hillslope analysis is in progress.

#### References

- Allison, R.J., 1990, Developments in a non-destructive method of determining rock strength: Earth Surface Processes and Landforms, v. 15, p. 571-577.
- Allison, R.J., 1991, Developments in a non-destructive method of determining rock strength: a reply: Earth Surface Processes and Landforms, v. 16, p. 473-476.
- Campbell, I.A., 1991, Comments on Allison's (1990) Developments in a non-destructive method of determining rock strength: Earth Surface Processes and Landforms, v. 16, p. 471-472.
- Day, M.J., 1980, Rock hardness: field assessment and geomorphic importance: Professional Geographer, v. 32, no. 1, p. 72-81.
- Grant, G.E., and Swanson, F.J., 1995, Morphology and processes of valley floors in mountain streams, Western Cascades, Oregon: *in* Costa, J.E., et al., eds., Natural and anthropomorphic influences in fluvial geomorphology, Geophysical Monograph 89, American Geophysical Union, p. 83-101.
- McCarroll, D., 1991, The Schmidt hammer, weathering, and rock surface roughness: Earth Surface Processes and Landforms, v. 16, p. 477-480.
- Melton, M. A., 1960, Intravalley variation in slope angles related to microclimate and erosional environment: Geological Society of America Bulletin, v. 71, p. 133-144.
- Montgomery, D.R., and Buffington, J.M., 1997, Channel-reach morphology in mountain drainage basins: Geological Society of America Bulletin, v. 109, no. 5, p. 596-611.
- Selby, M.J., 1993, Hillslope Materials and Processes, 2nd Edition: Oxford University Press, Oxford.
- Sklar, L.S, and Dietrich, W.E., 2001, Sediment and rock strength controls on river incision into bedrock: Geology, v. 29, no. 12, p. 1087-1090.
- Stewart, D.E., Lewis, R.S., and Link, P.K., 1995, 1996, 2001-2003, Unpublished mapping: Idaho Geological Survey.

## Why not Selby's (1980) Rock Mass Strength Index?

RMS is primarily a hillslope index.

It estimates the total strength of a mass of rock; this is useful for engineering applications, but "seldom relevant to geomorphic study" (Selby, 1980).

I am interested in fluvial processes where strength at a smaller scale is needed.

RMS index takes semi-quantitative measures and turns them back into a qualitative index!

# Variance analysis by lithology

| Groups         | Count              | Sum      | Average   | Variance |                |          |
|----------------|--------------------|----------|-----------|----------|----------------|----------|
| Yh/Yaq         | 246                | 402.5816 | 1.636511  | 0.019658 |                |          |
| Zdi-sy         | 722                | 1170.633 | 1.621376  | 0.024695 |                |          |
| Tdq/Tss        | 935                | 1521.702 | 1.627488  | 0.023102 |                |          |
| Tgd/Tgdf       | 170                | 287.6142 | 1.691848  | 0.012589 |                |          |
| ANOVA          | the State of State | and the  | to the ed |          |                |          |
| Source of      | the first          | 285      |           | 1.2.2.2  | Charles Martin | 6.4.5    |
| Variation      | SS                 | df       | MS        | F        | P-value        | F crit   |
| Between Groups | 0.714262           | 3        | 0.238087  | 10.63337 | 6.16E-07       | 2.609205 |
| Within Groups  | 46.3261            | 2069     | 0.022391  |          |                |          |
|                |                    |          |           |          |                |          |
| Total          | 47.04036           | 2072     |           | 1994     |                |          |